

FULL PAPER

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A long-term trend in the $F2$ -layer critical frequency as observed at Alma-Ata ionosonde station

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Abstract

In this study, we combine monthly median values for the $F2$ -layer critical frequency ($foF2$), measured at Alma-Ata ionosonde station [43.25°N, 76.92°E] between 1000 and 1400 (local time), with historical data on the monthly mean values for solar radio flux at 10.7 cm ($F10.7$) and the geomagnetic activity index (Ap) (available at <http://www.swpc.noaa.gov/>), over the period from 1957 to 2012. These data are used to derive long-term trends in the upper ionosphere and to discuss their importance in the context of coupling between solar and geomagnetic activity in the ionosphere at middle latitudes.

Keywords: Upper ionosphere; Long-term changes (trends); Solar-ionosphere interactions

Background

Since Rishbeth (1990) and Rishbeth and Roble (1992) predicted an effect on HF/VHF radio propagation due to the 'greenhouse cooling' associated with lowering of the $F2$ -layer, the long-term variations in the upper atmosphere's and ionosphere's parameters have been the subject of a number of observational and modeling studies (Mikhailov 2006; Lastovicka et al. 2008a; Danilov 2012; and references therein). These studies employed a variety of methods, approaches, and datasets to extract long-term trends in the $F2$ -layer critical frequency ($foF2$) and peak height ($hmF2$) for the ionospheric $F2$ -layer, and the majority were conducted using ionosonde observations. Table 1 summarizes these studies.

Potential drivers of long-term trends in $foF2$ are widely discussed by Yue et al. (2006), Lastovicka (2009), Danilov (2012), and references therein and include long-term variations in solar and geomagnetic activity, increasing concentrations of greenhouse gases (e.g., CO_2 , CH_4) and anthropogenic changes to the ozone layer and the distribution of water vapor. Our study focuses on the role of solar and geomagnetic activity in long-term $foF2$ trends, using $foF2$ data routinely measured over Kazakhstan at the Alma-Ata ionosonde station [43.25°N, 76.92°E]. Data used in this study cover about five solar cycles between 1957 and 2012. Data measured at the Alma-Ata station

between 1958 and 1994 have already been used to derive long-term trends, independent of geomagnetic activity (e.g., Danilov 2003). However, this study is the first to use the extended dataset up to the year 2012 to derive long-term trends in $foF2$ (we assume the trend is a long-term linear change in $foF2$ over the period between 1957 and 2012).

Methods

For this study, we used monthly median values of $foF2$ averaged over 5 h, between 1000 and 1400 (local time) (Figure 1a). Monthly mean values of the solar radio flux at 10.7 cm ($F10.7$) and geomagnetic activity index (Ap) (available at <http://www.swpc.noaa.gov/>) were also studied as the characteristics of solar and geomagnetic activities can strongly affect the ionosphere (Figure 1b,c). As expected, Figure 1 shows that the temporal variations in all three parameters are dominated by changes in solar activity related to the solar cycle. Assuming a second-order polynomial dependence on $F10.7$, $foF2'$ is defined as follows (see the regression line in Figure 2a):

$$foF2' = a + b \times (F10.7) + c \times (F10.7)^2 \quad (1)$$

A higher-order (cubic) regression, as used by Chen et al. (2014), does not provide any significant improvement to the fit: $R^2 = 0.810324$ for a second-order regression versus $R^2 = 0.810385$ for a third-order regression). The regression defined in Equation 1 was then used to remove variations in $foF2$ related to the solar activity effect, allowing monthly

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Table 1 List of authors who derived long-term trends in foF2 and their findings

Authors	Site: name or number (N) of stations	Years of observation	Trends (MHz y ⁻¹)
Danilov (2002)	Sverdlovsk [56.43°N, 58.57°E]	1948-1994	-0.015
Danilov (2003)	N = 21, φ > 30°	1958-mid-nineties	-0.012
Danilov (2013)	N = 12	1990-2010	-0.024 (1400 local time), -0.054 (after sunset)
Mielich and Bremer (2013)	N = 124	1948-2009	-0.003 to 0.0038
Yue et al. (2006)	N = 19, [42.9°S to 62.0°N, ca. 130°E]	1948-2005	-0.005
Lastovicka et al. (2006, 2008b)	Juliusruh [54.6°N, 13.4°E]	1976-1996	-0.01 to -0.02
Ghabahou et al. (2013)	Ouagadougou [12.4°N, 358.5°E]	1966-1998	-0.015
Khaitov et al. (2012)	Tomsk [56.5°N, 84.9°E]	1937-2011	-0.008 to -0.014
Mikhailov (2006)	Slough [51.48°N, -0.57°E]	1935-2000	-0.00086

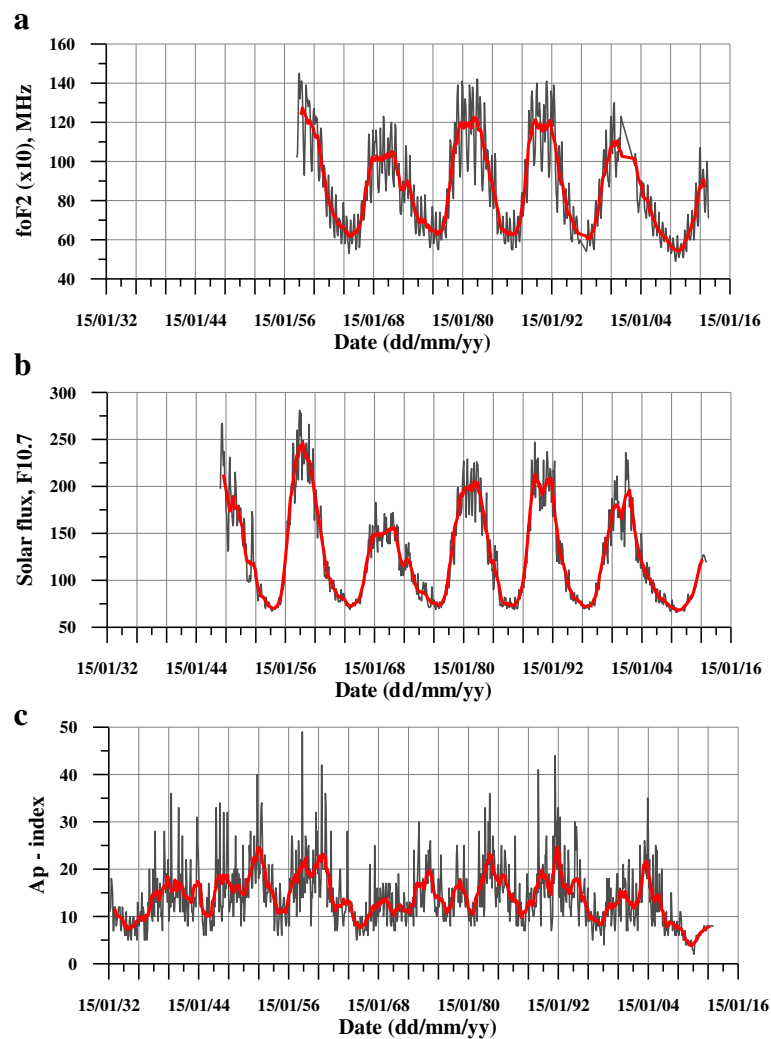
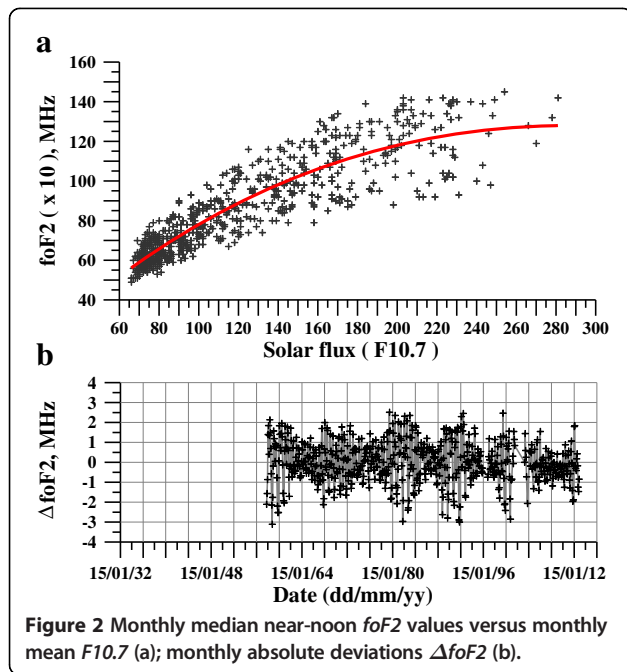
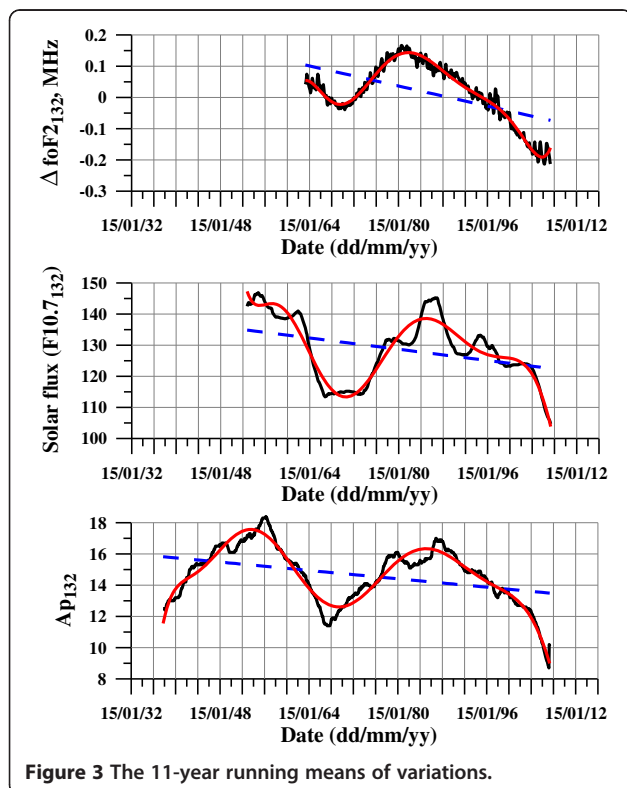


Figure 1 Monthly median near-noon foF2 (a), monthly mean F10.7 (b) and Ap (c) values.



absolute deviations (defined as $\Delta foF2 = foF2 - foF2$), which can potentially reveal long-term trends in $foF2$, to be calculated (Figure 2b). However, Figure 2b shows that the correlation between $foF2$ and $F10.7$ determined via Equation 1 only accounts for around 80% ($R^2 = 0.810324$), of the variations in $foF2$ and the majority of the $\Delta foF2$



variability linked to the 11-year solar cycle. Therefore, to obtain an independent picture of long-term trends in the upper ionosphere, the 11-year (132 months) running mean values of the monthly absolute deviations ($\Delta foF2_{132}$) were calculated over the entire dataset according to the method set out by Mikhailov (2006). This 11-year smoothing technique was also applied to the $F10.7$ and A_p datasets but reduced the available period for study to between 1962 and 2006.

Results and discussion

Figure 3 presents the calculated values for $\Delta foF2_{132}$, $F10.7_{132}$, and A_p_{132} . Figure 3 shows a pronounced similarity in the $\Delta foF2_{132}$, $F10.7_{132}$, and A_p_{132} variations, which display negative correlations with time and a repeating pattern with a period of ca. 30-32 years. This implies that geomagnetic activity and the Earth's ionosphere are strongly controlled by solar activity. However, it should be noted that the negative trend in geomagnetic activity found here contradicts the generally accepted increase in geomagnetic activity observed throughout the twentieth century (e.g., Clilverd et al. 1998). However, in detail, the geomagnetic activity increased throughout the first half of the twentieth century (along with solar activity) then stabilized (with some increase in A_p seen at the end of the 1950s), and then decreased until the beginning of the twenty-first century (with another smaller A_p peak observed in the 1980s) (Figures 3 and 4). This study spans the interval from 1957 to 2012, and our data match the overall decrease in A_p observed over these years. The trend with a period of ca. 30-32 years is likely to have a solar origin, as it matches a period of 31.1 years that has been found elsewhere in sunspot number spectral analyses (Echer et al. 2004; Clúa de Gonzalez et al. 1993). It has also been suggested that this period of 31 years is the origin of the 35-year Brückner climatic periodicity (Raspopov et al. 2000). Figure 3 shows evidence of the same solar periodicity in $foF2$ long-term variations. The Fisher (F) parameter for $foF2$ data confirmed that the clear negative trend (ca. $-0.0038 \text{ MHz y}^{-1}$) was significant with a confidence level of 95%-99%.

Using a similar method to that described for $foF2$ above, we calculated the dependence of monthly mean A_p on $F10.7$, allowing the variations in A_p related to the solar cycle to be clearly seen (Figure 5). Assuming a linear dependence, we defined $(A_p)'$ as a function of $F10.7$ and obtained absolute deviations (ΔA_p) as follows:

$$(A_p)' = a + b \times F10.7 \quad (2)$$

$$\Delta A_p = A_p - (A_p)' \quad (3)$$

Figure 5 displays the observed A_p values (black crosses) versus $F10.7$, together with the linear regression

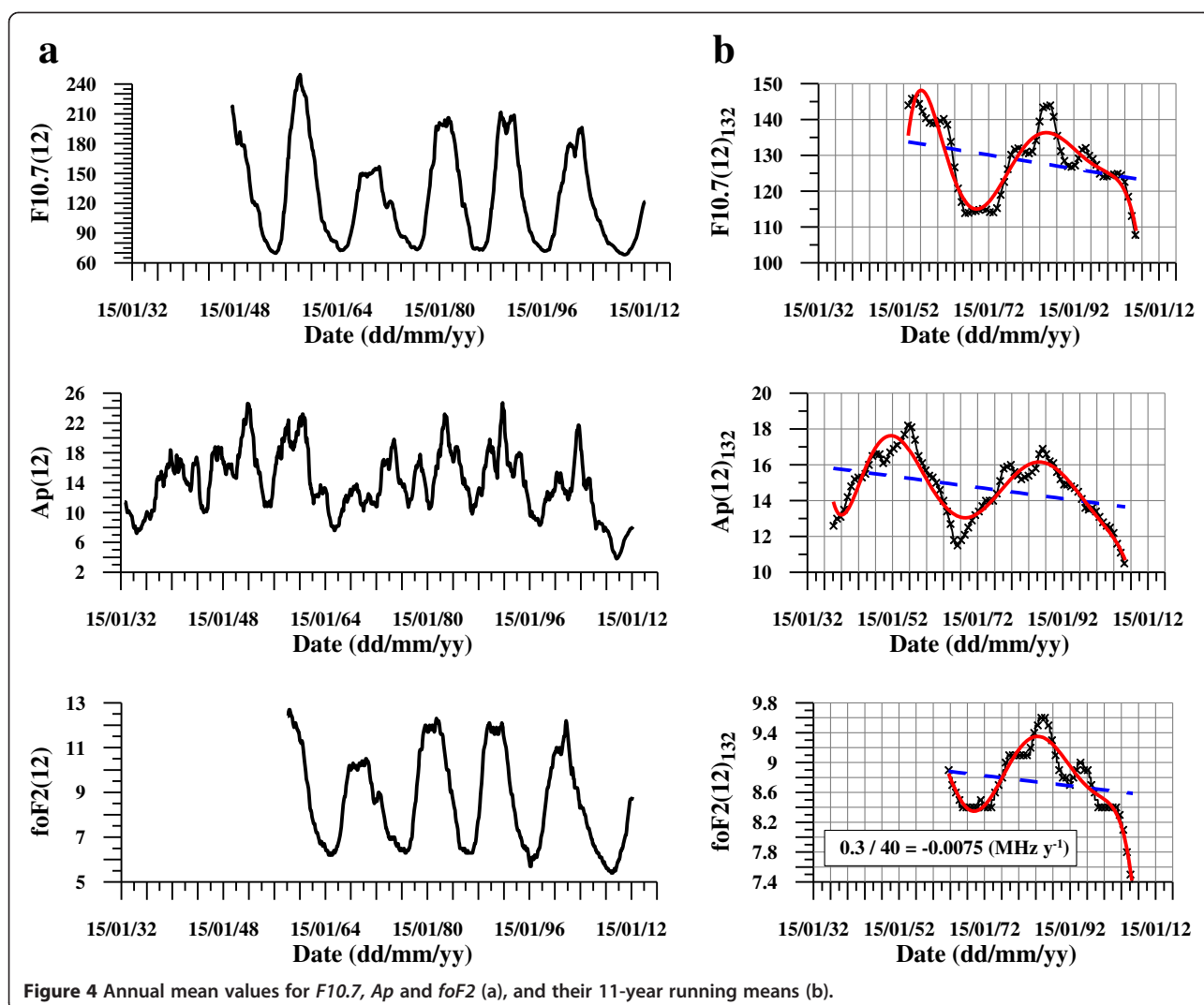
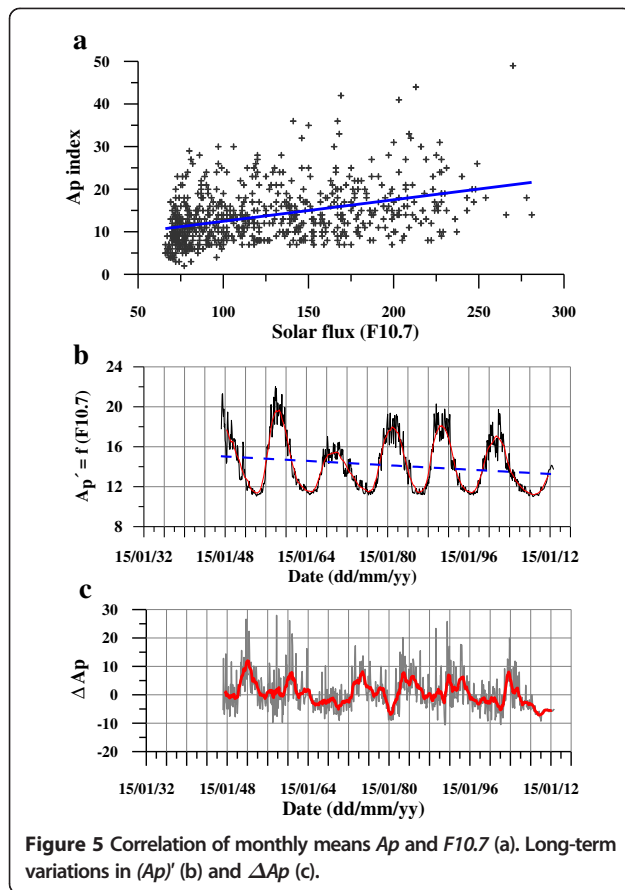


Figure 4 Annual mean values for $F10.7$, A_p and $foF2$ (a), and their 11-year running means (b).

line (solid line), the variations in A_p related to solar forcing, and the variations in ΔA_p with time. Approximately 16% of the variations in the geomagnetic field can be explained by the linear relationship between geomagnetic and solar activities ($R^2 = 0.15798$, Figure 5a) and the majority variations in ΔA_p are linked to the 11-year solar cycle. Peaks in ΔA_p are slightly shifted (by about 2 to 3 years) relative to the falling phase of the 11-year solar cycle (Figure 5c). Taking this shift into account for the regression calculation did not result in a significantly better fit ($R^2 = 0.1698$). These results show that the geomagnetic activity (described by A_p) is strongly linked to the solar cycle phase (solar activity is described by $F10.7$) and in this study, we were unable to exclude variations in $foF2$ related to geomagnetic activity. Analyzing geomagnetic data observed at Kakioka (Japan) and Gnamara (Australia) over almost five solar cycles, Yamazaki and Yumoto (2012) recently found that solar activity controls not only the stationary component

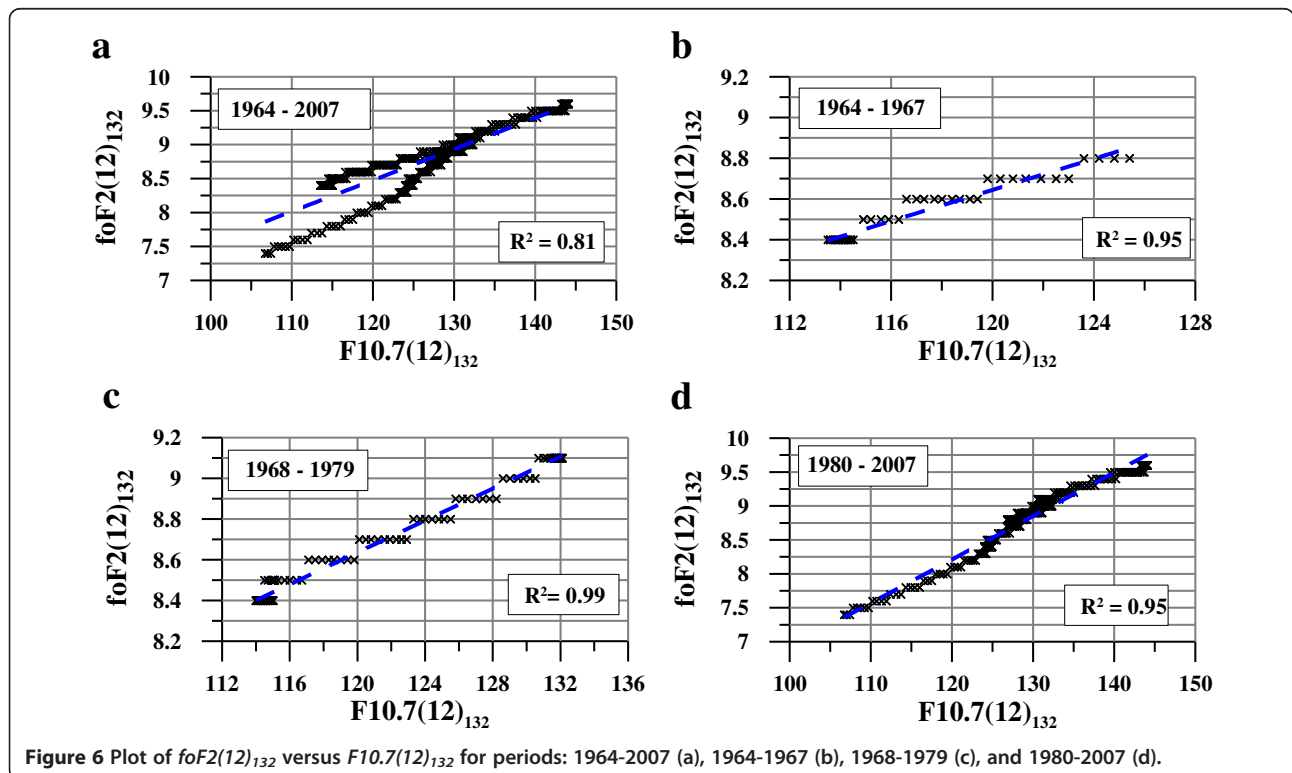
of the geomagnetic solar quiet daily variation field (S_q) but also the annual and semi-annual components. They report that all three components have a positive linear correlation with sunspot numbers. Thus, the positive linear correlation between A_p and $F10.7$ found in this study confirms Yamazaki and Yumoto's findings and shows the existence of a long-term coupling between solar and geomagnetic activity that could be used to further our understanding of solar-terrestrial relations.

Our results also show that $foF2$ strongly depends on solar activity and shows a negative temporal trend between 1957 and 2012 (about $-0.0038 \text{ MHz } y^{-1}$), although the magnitude of this trend is probably too small value to be of practical use. However, it should be noted that the sign of the deduced trend can be dependent on choice of time period for trend analysis. Periods of increasing solar activity (1970-1984) are seen to correspond to positive trends in $foF2$ and periods of decreasing solar activity (1956-1968, 1986-2004) to



negative trends in $foF2$ (Figure 3). Therefore, periods of several solar-cycle observations should be used to obtain reliable trend estimates from the data series.

In addition to the material presented above, we derived a picture of long-term changes in the upper ionosphere using annual mean values for Ap and $F10.7$ ($Ap(12)$ and $F10.7(12)$) and annual median values for $foF2$ ($foF2(12)$). Following a similar method to that described above, Figure 4a,b shows the variations in the 11-year running means $foF2(12)_{132}$, $F10.7(12)_{132}$, and $Ap(12)_{132}$ for the analyzed period. Figure 4b shows that long-term trends are similar to those seen in Figure 3, which supports our conclusion that variations in Ap and $foF2$ are dominantly affected by solar cycles as represented by $F10.7$. One exception to this conclusion is the somewhat higher $foF2$ trend ($-0.0075 \text{ MHz y}^{-1}$) than that found using the regression method and including an $F10.7$ correction ($-0.0038 \text{ MHz y}^{-1}$). Table 1 shows that the higher $foF2$ trend is close to those calculated by Danilov (2002, 2003), Lastovicka et al. (2006, 2008b), Khaïtov et al. (2012), and Ghabahou et al. (2013), whereas the weaker $foF2$ trend more closely matches that calculated by Mielich and Bremer (2013). Here, we can only note that twice removing the solar element of variations in $foF2$ (using the regression method and the 11-year running mean) provides a weaker $foF2$ trend than that obtained using only the 11-year smoothing.



Additionally, we calculated regressions for $foF2(12)_{132}$ as a function of $F10.7(12)_{132}$ for different periods. Figure 6a shows that for the total interval (07.1963 to 08.2006), the correlation between the two variables forked into two distinct point groups assuming different relationships between $foF2(12)_{132}$ and $F10.7(12)_{132}$ for different phases of the ca. 32-year cycle, a coefficient of determination (R^2) of 0.81 was obtained. Figure 6b,d shows linear relationships for the intervals showing a decrease (1964-1967; 1980-2007) and increase (1968-1979) in solar activity. It was found that 95% and 99% of the variations in $foF2(12)_{132}$ could be explained by linear relationships between $foF2(12)_{132}$ and $F10.7(12)_{132}$ for the decreasing and increasing intervals, respectively. The remaining variations in $foF2(12)_{132}$ are not explained by solar activity.

Conclusions

In this study, we derived a picture of long-term trends in $foF2$ for the ionosphere, using data from the mid-latitude ionosonde station at Alma-Ata [43.25°N, 76.92°E] observed over about five solar cycles between 1957 and 2012. We showed that solar activity (as represented by $F10.7$) is significantly correlated with variations in $foF2$ and A_p . In addition to the well-known 11-year solar cycle, the Sun also exhibits a cycle of about 30-32 years, which matches the period of trends observed in A_p and $foF2$. A negative trend is seen in long-term variations in $foF2$ between 1957 and 2012, and the magnitude of this trend was found to be -0.0038 and -0.0075 MHz y^{-1} for monthly absolute deviations ($\Delta foF2$) and annual mean median $foF2$, respectively. This trend is considered too small to have practical meaning. It was found that 95% and 99% of the total variation in $foF2(12)_{132}$ could be explained by linear relationships between $foF2(12)_{132}$ and $F10.7(12)_{132}$ for periods increasing and decreasing solar activity, respectively. The remaining variations in $foF2(12)_{132}$ cannot be explained by solar activity.

Competing interests

The authors declare that they have no competing interests.

Authors' contributions

All authors have been involved in data interpretation, drafting the manuscript, and revising it critically. All authors read and approved the final manuscript.

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