# Mechanisms of hydrogen generation during the mechanochemical treatment of biotite within D<sub>2</sub>O media

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(Received May 31, 2004; Revised September 13, 2004; Accepted November 16, 2004)

The mechanism of hydrogen generation during the mechanochemical treatment of biotite was examined by grinding experiments using a ball mil in H<sub>2</sub>O or D<sub>2</sub>O as a grinding media. From the linear relationship between the amount of generated hydrogen and the increase of the surface area of ground powders, the hydrogen productivity of biotite is estimated to be  $0.036 \ \mu \text{mol/m}^2$ , which agrees with the previous results in spite of the difference in the grinding conditions. D<sub>2</sub> analyses by a mass spectrometry indicate that the produced amount of D<sub>2</sub> accounts for only 10% of the total hydrogen and that more than 90% of hydrogen takes a form of a mixture of HD and H<sub>2</sub>. The observed isotope distribution clearly indicates that hydroxyls within the crystal structure can be a major source for the generation of hydrogen. Hydrogen generation originated from hydroxyls may indicate the higher hydrogen productivity of phyllosilicates than those of quartz and alkali feldspar. **Key words:** Hydrogen generation, mechanochemical treatment, biotite, hydrogen isotope.

# 1. Introduction

There are many reports on hydrogen anomaly associated with seismic events (e.g. Wakita *et al.*, 1980; Sugisaki *et al.*, 1983; Sugisaki and Sugiura, 1986; Ito *et al.*, 1999). Hydrogen observed along active faults is thought to be generated by mechanochemical reactions within the fault zones (Wakita *et al.*, 1980). This unique origin of hydrogen has attracted extensive interests. Since hydrogen is produced by chemical reactions where fault rocks are destructed, it is expected that hydrogen may carry valuable information concerning the seismogenic regions (Koizumi, 1997).

Hydrogen generation by mechanochemical reactions between a crushed rock surface and water has been confirmed by means of crushing experiments using quartz and granite powder in the presence of water vapor (Kita et al., 1982). Sugisaki et al. (1983) also reacted several kinds of powdered samples with water in a flask and detected molecular hydrogen generation. According to the recent grinding experiments using a ball mill by the present authors (Kameda et al., 2003), hydrogen productivity for some single crystals was evaluated by comparing the amount of generated hydrogen and the freshly created surface area of ground samples. It is reported that phyllosilicates such as biotite and muscovite can effectively produce hydrogen and that more than 60% of hydrogen can be attributed to biotite and chlorite in the case of granite. Thus, it is important to interpret the high hydrogen productivity of phyllosilicates in order to understand the natural occurrences and origin of hydrogen along active faults. Kameda *et al.* (2003) suggested that the high productivity seems to be correlated with the existence of hydroxyls within the crystal structures. If the hydroxyls actually cause hydrogen production, the observed hydrogen can be thought to have two origins: one is hydrogen atoms from the grinding liquid media (water molecules) and the other from the hydroxyls in a crystal.

In this study, we have conducted the wet grinding experiments of biotite within  $D_2O$  as a grinding liquid media in order to elucidate the role of hydroxyls for hydrogen generation. The analyses of stable isotopes for the generated hydrogen may provide a clue to the origin of hydrogen and its generation mechanisms.

# 2. Experimental

#### 2.1 Milling of the powders and analysis

Grinding was conducted by using a rotational ball-mill (Fritsch P-6), zirconia mill pot (80 cc) and zirconia beads (10 mm $\phi \times 20$ ) at room temperature (25°C) under atmospheric pressure. Processed samples were prepared from the single plate-like crystals of biotite (about 10 cm  $\times$  10  $cm \times 0.5 cm$ ). The chemical compositions (in weight percentage) were determined by electron-probe microanalyzer (EPMA) and Fe (II) and Fe (III) contents measured by wet analyses: Na<sub>2</sub>O, 0.153; Al<sub>2</sub>O<sub>3</sub>, 17.93; FeO, 19.26; Fe<sub>2</sub>O<sub>3</sub>, 0.26; K<sub>2</sub>O, 8.397; MgO, 8.576; SiO<sub>2</sub>, 36.818; TiO<sub>2</sub>, 2.384; MnO, 0.056; Cr<sub>2</sub>O<sub>3</sub>, 0.002 (±0.5% accuracy). A single crystal of biotite was gently crushed by a tungsten-carbide motor and washed ultrasonically in ethanol for three times. A dried powder was sieved to grain sizes ranging from 0.15 to 0.512 mm. A starting powder was preheated at 150°C for more than 2 hours to remove adsorbed surface water.

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Fig. 1. The relationship between the partial pressures of  $m/e = 4 (D_2^+)$  and  $m/e = 20 (D_2O^+)$ .

3 g of biotite powder and 10 cc pure of either H<sub>2</sub>O (Millipore water;  $15M\Omega$  or D<sub>2</sub>O (Wako chemicals; 99.5% purity) were put in a mill pot within a glove box (Unico UN-650F) in order to exchange an internal gas phase by Ar (99.995% purity) for H<sub>2</sub>O experiments and by N<sub>2</sub> (99.9995% purity) for D<sub>2</sub>O experiments, respectively. O<sub>2</sub> content within the glove box was estimated to be <100 ppm as was monitored by an O<sub>2</sub> meter (Toray LC-750). Duration of milling was 15, 30, 45 and 60 minutes at 500 rpm. A gas phase after grinding was sampled by gas-tight syringe (Ito MS-GAN 250). The total hydrogen content was analyzed by gas chromatography (Yanaco G-2700; MS-5A column) equipped with a thermal conductivity detector (TCD). D<sub>2</sub> content was analyzed by mass spectrometer as described in the succeeding section. The ground sample was separated by centrifugation at 4500 rpm for 20 min and the surface area of a ground powder was measured by the BET (Brunauer-Emmet-Teller) method (Beckman Coulter SA-3100) at liquid nitrogen temperature (77 K).

# 2.2 Mass spectrum

A generated amount of D<sub>2</sub> in a N<sub>2</sub> matrix was measured by quadrupole mass spectrometer (QMS, HIDEN HAL-201). For the measurements, 1 cc of gas sample was injected into the vacuum line ( $\sim 1.0 \times 10^{0}$  Pa) attached to the QMS. A sample was thereafter carried into the analyzer  $(\sim 1.0 \times 10^{-7} \text{ Pa})$  through a flow control value. During measurements, gas pressures in the line and analyzer were kept around 5.5  $\times$  10<sup>3</sup> Pa and 7.3  $\times$  10<sup>-5</sup> Pa, respectively. An electron energy of 80 eV and an emission current of 20  $\mu$ A were used for analyses. Seven scans were averaged to obtain a peak area. Because there was D<sub>2</sub>O vapor in the gas sample, a  $D_2^+$  ion (m/e = 4) was always detected as a blank due to the fragmentation of D<sub>2</sub>O molecules. In order to evaluate the effect of D<sub>2</sub> originated from D<sub>2</sub>O, a D<sub>2</sub>O gas sample without D2 was measured at various concentrations for determining the peak area ratio between m/e = 20 $(D_2O^+)$  and m/e = 4  $(D_2^+)$ . Figure 1 shows the linear correlation between the peak area of  $D_2^+$  ( $[D_2^+]_{blank}$ ) and that of  $D_2O^+$  ([ $D_2O^+$ ]), leading to the following regression line:

$$[D_2^+]_{blank} = 0.0032 \times [D_2O^+]$$
 (1)

The amount of original  $D_2$  ( $[D_2^+]$ ) was obtained by subtracting the area of blank  $D_2^+$  ( $[D_2^+]_{blank}$ ) from the total area of  $D_2^+$  ( $[D_2^+]_{total}$ ):

$$[D_2^+] = [D_2^+]_{\text{total}} - [D_2^+]_{\text{blank}}$$
(2)

Neglecting the change of N<sub>2</sub> partial pressure during experiments,  $[D_2^+]$  can be corrected by the area of m/e = 14 (N<sup>+</sup>) as a reference, which was produced by the fragmentation of N<sub>2</sub>. From a calibration curve using a standard D<sub>2</sub> gas, the peak area was converted to the D<sub>2</sub> concentration. Unfortunately, because of the high blank of m/e = 2 (H<sub>2</sub><sup>+</sup>) and m/e = 3 (DH<sup>+</sup>) caused by fragmentation of D<sub>2</sub>O and H<sub>2</sub>O in the analyzer, the amounts of DH and H<sub>2</sub> could not be determined successfully in this analytical method.

# 3. Results and Discussion

#### 3.1 Wet grinding within H<sub>2</sub>O media

Table 1 shows the results of H<sub>2</sub> measurements and surface area analyses after grinding biotite in H<sub>2</sub>O liquid media. Both the H<sub>2</sub> concentration and the specific surface area monotonically increase with the duration of grinding. After 60 minutes, the H<sub>2</sub> concentration reaches 168 ppm, which is smaller than 648 ppm at the same duration observed in a previous study (Kameda et al., 2003). The specific surface area after 60 minutes grinding is determined to be 4.42 m<sup>2</sup>/g, which is also lower than the results obtained by Kameda et al. (2003). Such lower values seem to be caused by the smaller amounts of samples, fewer balls and slower milling conditions. Comparing these results, the relationship between the increase of the total surface area of ground powders and the generated amount of H<sub>2</sub> in molar quantity can be obtained (Fig. 2). The amount of H<sub>2</sub> is linearly correlated with the increase of a total surface area with the slope of 0.036  $\mu$ mol/m<sup>2</sup>. Although the grinding conditions are different from those by Kameda et al. (2003), the H<sub>2</sub> productivity fairly agrees with each other.

Grinding time	H <sub>2</sub> concentration		D <sub>2</sub> concentration		Specific surface area
(min)	(ppm)	$(\mu mol)$	(ppm)	$(\mu mol)$	$(m^2/g)$
0	_	—	_	_	0.691
15	50.3	0.12	2.4	0.0072	2.09
30	84.5	0.22	5.5	0.0165	2.86
45	135	0.35	7.4	0.0222	3.9
60	168	0.43	12.7	0.0381	4.42

Table 1. The experimental results of  $H_2$  and  $D_2$  analysis. The specific surface areas of the ground biotite are also shown.



Fig. 2. The linear relationship between the amount of generated  $H_2$  ( $\mu$ mol) and the increase of the surface area of ground biotite (m<sup>2</sup>). The values in the parentheses indicate the  $H_2$  productivity of biotite.

As presented previously in Kameda *et al.* (2003), the hydrogen productivity of biotite is much higher than those of quartz  $(1.0 \times 10^{-3} \ \mu \text{mol/m}^2)$  and alkali feldspar  $(3.9 \times 10^{-3} \ \mu \text{mol/m}^2)$ . Hydrogen generation during grinding of quartz or alkali feldspar can be explained by the chemical reactions between the mechanoradicals on a fresh surface of minerals and grinding liquid media as follows (Schrader *et al.*, 1969; Kita *et al.*, 1982):

$$2(\equiv Si\bullet) + 2H_2O \rightarrow 2(\equiv SiOH) + H_2$$
(3)

Higher hydrogen productivity of biotite is likely related to the existence of hydroxyls in its structure (Kameda *et al.*, 2003). The prospect was supported by dry grinding experiments using kaolinite (Kameda *et al.*, 2004). Hydrogen generation during the grinding of kaolinite was observed even under dry conditions. Furthermore, infrared (IR) absorption measurements reveal that the intensity due to hydroxyls decreases in accordance with the hydrogen generation. If this is the case, the observed hydrogen during grinding biotite in H<sub>2</sub>O liquid media may be produced by hydrogen atoms from grinding media and/or from hydroxyl groups.

# 3.2 Wet grinding within D<sub>2</sub>O media

For the quantitative discussion on the origin of hydrogen atoms,  $D_2O$  liquid media were used for grinding procedures. In this case, we can suppose that the generated hydrogen may be composed of a mixture of  $H_2$ , HD and  $D_2$ . At first, the apparent hydrogen concentration after 60 minutes grinding in  $D_2O$  was analyzed by gas chromatography. Assuming all of hydrogen atoms were from  $H_2$ , 154 ppm of hydrogen was detected and this seems a little lower than the result in  $H_2O$  experiments (168 ppm). Since the thermal conductivity detector (TCD) was used for GC measurements, the slightly lower concentration may suggest that  $D_2$ or DH whose thermal conductivity is lower than that of  $H_2$ was mixed in the apparent hydrogen peak.

Because it is difficult to separate a mixture of hydrogen (H<sub>2</sub>, HD and D<sub>2</sub>) into each component by the present GC measurements, mass spectrometry was adopted to determine the amount of generated D<sub>2</sub>. Figure 3 shows the relationship between the D<sub>2</sub> concentration and the grinding time. The concentration increased with the duration of grinding time until 12.7 ppm of D<sub>2</sub> was detected after 60 minutes. Surprisingly, this value indicates that only ~10% of total hydrogen was generated by D<sub>2</sub>O grinding media according to the following equation:

$$2(\equiv Si\bullet) + 2D_2O \rightarrow 2(\equiv SiOD) + D_2 \tag{4}$$

To obtain the  $D_2$  productivity, the generated amount of  $D_2$  ( $\mu$ mol) was compared with the increase of the total surface area (Fig. 4). Although the points are scattered, the slope of the regression line between them indicates that the  $D_2$  productivity is within 0.002–0.0035  $\mu$ mol/m<sup>2</sup>. Although the structural change due to the destruction of biotite lay-



Fig. 3. The relationship between the amount of the generated hydrogen concentration (ppm) and the increase of the specific surface area of ground biotite (m<sup>2</sup>).



Fig. 4. The increase of the amount of  $D_2$  with the total surface area of ground biotite.

ers may be different from that of quartz and K-feldspar, the amount of generated hydrogen from grinding media is comparable with those of quartz and K-feldspar.

On the other hand, more than 90% of the observed hydrogen can be attributed to HD and H<sub>2</sub> (Fig. 3) and the existence of such large amounts of HD and H<sub>2</sub> indicates that hydroxyls are considered to be an important source for the hydrogen formation. Because the surface of biotite was always in contact with D<sub>2</sub>O media during milling, H atoms as a precursor for H<sub>2</sub> or HD molecules should be formed within the crystal structure. One possible mechanism is the reduction of a proton by Fe (II) in a biotite crystal:

$$H^+ + Fe(II) \rightarrow H \bullet + Fe(III)$$
 (5)

Miller and Oulton (1970) observed a collapse of hydroxyl groups of kaolinite on IR spectra during the percussive milling and explained it by a *prototropy* effect, i.e. the transfer of protons within a kaolinite structure to form  $H_2O$  molecules. Such dehydroxylation (mechanochemical dehydroxylation) accompanied with the crystal destruction was also reported for montmorillonite (Blahoslav and

Gunther, 1981) and pyrophyllite (Sánchez-Soto and Pérez-Rodríguez, 1989). It can be supposed that prototropy effects also proceed within a biotite crystal during grinding in this study. If proton transfer does occur, a Fe (II) ion possibly reduces the proton to form a H atom. Since the FeO content of biotite is 19.26% as given in the experimental section, the molar quantity of Fe (II) corresponds to  $7.9 \times 10^{-3}$  mol for 3 g sample, which is sufficiently large (about two orders) to produce the observed H<sub>2</sub>. However, Kameda et al. (2003) observed that muscovite with the about 1/6 smaller content of Fe (II) than biotite (unpublished data) also exhibits the higher H<sub>2</sub> productivity (0.022  $\mu$ mol/m<sup>2</sup>). This inconsistency between the Fe (II) content and the amount of generated H<sub>2</sub> may suggest that not all H• radicals are produced by proton reduction. This leads us to the postulation of another possible mechanism or the direct interaction between Si• and Al-OH in the octahedral layer of biotite within the crystal as below.

$$(\equiv Si\bullet) + AlOH \rightarrow \equiv Si - O - Al + H\bullet$$
(6)

This reaction seems to be reasonable because it can explain

the high hydrogen productivity of muscovite, although further studies should be performed to validate this model. In any cases, it is postulated that the mechanochemical treatment of biotite causes the destruction of a crystal structure to lead to the reaction of the hydroxyls successively producing a H atom to form hydrogen.

# 4. Conclusions

Pure biotite powders were ground within  $H_2O$  or  $D_2O$  as a liquid media. The  $H_2$  productivity of biotite obtained in the present experiments agrees with that of the previous study in spite of the difference in the grinding conditions.  $D_2$  analyses revealed that more than 90% of hydrogen was composed of HD and  $H_2$ , suggesting that the hydroxyls in a layer are a major source for the generation of hydrogen during the mechanochemcal treatment of biotite. Furthermore, it can be pointed out that the present results are principal for studying the stable isotopic analyses of  $H_2$  along active faults to discuss the reaction temperature or the origin of  $H_2$ in a future work.

Acknowledgments. This work was supported by a Research Fellowship from the Japan Society for the promotion of Science for Young Scientists, the Grants-in-Aid for Scientific Research System of the Japanese Society for Promotion of Science (13440134) and by a grant for the comprehensive research program on the flow and slip processes in and below seismogenic regions, sponsored by the Ministry of Education, Culture, Sports, Science, and Technology, Japan.

# References

- Blahoslav, C. and K. Gunther, Mechanism of montmorillonite structure degradation by percussive grinding, *Clay Miner.*, 16, 151–162, 1981.
- Ito, T., K. Nagamine, K. Yamamoto, M. Adachi, and I. Kawabe, Preseismic hydrogen gas anomalies caused by stress-corrosion process preceding earthquakes, *Geophys. Res. Lett.*, 26, 2009–2012, 1999.
- Kameda, J., K. Saruwatari, and H. Tanaka, H<sub>2</sub> generation in wet grinding of granite and single-crystal powders and its implications for H<sub>2</sub> concentration on active faults, *Geophys. Res. Lett.*, doi:10.1029, 2003.
- Kameda, J., K. Saruwatari, and H. Tanaka, H<sub>2</sub> generation during dry grinding of kaolinite, J. Colloid. Interface. Sci., 275, 225–228, 2004.
- Kita, I., S. Matsuo, and H. Wakita,  $H_2$  generation by reaction between  $H_2O$  and crushed rock: An experimental study on  $H_2$  degassing from the active fault zone, *J. Geophys. Res.*, **87**, 10,789–10,795, 1982.
- Koizumi, N., Review of geochemical research for earthquake prediction, *J. JSNDS*, 16-1, 41–60, 1997 (Japanese with English abstract).
- Miller, J. G. and T. D. Oulton, Prototropy in kaolinite during percussive grinding, *Clays Clay Miner.*, 18, 313–323, 1970.
- Sánchez-Soto, P. J. and J. L. Pérez-Rodríguez, Formation of mullite from pyrophyllite by mechanical and thermal treatments, *J. Am. Ceram. Soc.*, 72, 154–157, 1989.
- Schrader, R., R. Wissing, and H. Kubsch, Zur Oberflachenchemie von mechanisch activiertem Quartz, Z. Anorg. Allg. Chem., 365, 191–198, 1969.
- Sugisaki, R. and T. Sugiura, Gas anomalies at three mineral springs and a fumarole before an inland earthquake, central Japan, *J. Geophys. Res.*, 91, 12,296–12,304, 1986.
- Sugisaki, R., M. Ido, H. Takeda, Y. Isobe, Y. Hayashi, N. Nakamura, H. Satake, and Y. Mizutani, Origin of hydrogen and carbon dioxide in fault gases and its relation to fault activity, *J. Geol.*, **91**, 239–258, 1983.
- Wakita, H., Y. Nakamura, I. Kita., N. Fujii, and N. Notsu, Hydrogen release: New indicator of fault activity, *Science*, 210, 188–190, 1980.
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