

Low-temperature crystallization of thin silicate layer on crystalline Fe dust

Chihiro Kaito, Akihito Kumamoto, Yoshio Saito, and Ryoichi Ono

Laboratory for Nano-Structure Science, Department of Physics, Ritsumeikan University,
1-1-1 Nojihigashi, Kusatsu-shi, Shiga 525-8577, Japan

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The crystallization of an amorphous SiO layer covering Fe crystal grains has been clarified by high-resolution transmission electron microscopic (HRTEM) observation. Cristobalite crystals were produced preferentially on the (110) surface of Fe particles by the oxidation of silicon crystallites in the SiO layer, i.e. the oxidation energy of the silicon crystallites resulted in the epitaxial growth of the oxide layer on the Fe surface. The chemical reaction energy due to the oxidation of silicon crystallites in the SiO layer was concentrated at the interface of the crystal and the amorphous layer. Crystal growth took place from the Fe grain surface.

Key words: Low-temperature crystallization, silicate layer, crystalline dust, cristobalite, oxidation energy.

1. Introduction

To explain how cometary silicates crystallize under pressure in volatile interstellar ice in their parent comets, we experimentally demonstrated the possibility of a chemical reaction-driven crystallization (Kaito *et al.*, 2007), which is called nonthermal crystallization (Yamamoto and Chigai, 2005; Yamamoto *et al.*, 2010), using laboratory-synthesized amorphous Mg-bearing silicate grains (Kamitsuji *et al.*, 2005). If the crystalline grain surface is covered with an amorphous silicate layer, the crystallization of the amorphous silicate layer may be different due to the different rates of thermal diffusion between amorphous and crystal layers, as observed during the crystallization of a carbon layer covering Pt clusters (Shintaku *et al.*, 2006; Kaito *et al.*, 2007). In a previous successful low-temperature crystallization of amorphous dust, iron grains covered with a SiO layer were used as the model sample, as schematically shown in Fig. 1.

In a previous study (Kaito and Shimizu, 1984), a vacuum-deposited SiO film was shown to be composed of Si and SiO₂ (α -cristobalite) crystallites about 1 nm in diameter. By heating the film in air at various temperatures up to 500°C, the film changed from orange to being transparent while maintaining an amorphous structure with a halo diffraction pattern. Intensity analyses of the electron diffraction (ED) pattern and infrared spectroscopy were carried out to evaluate the structure alteration (Morioka *et al.*, 1998). The mixed film of Si and SiO₂ altered from α -cristobalite to β -cristobalite (250°C) and β -quartz (500°C) accompanied by the oxidation of Si crystallites in air. Infrared spectroscopy was carried out to evaluate the structure alteration (Morioka *et al.*, 1998). In this paper, Fe crystal grains covered with SiO film were heated at 100°C in the stable region for α -

cristobalites in air. The alteration of the SiO film was detected by HRTEM and is discussed as a low-temperature crystallization effect due to the oxidation energy of silicon crystallites in the SiO film.

2. Experimental

Fe wire (0.5-mm ϕ) was wound on the tungsten wire with a 0.5 mm ϕ filament, which had been preheated in a vacuum of 10⁻⁵ Torr. The chamber was a glass cylinder (17 cm diameter and 30 cm height) covered with stainless steel on top and connected to a high-vacuum exhaust through a valve at its bottom. After evacuation of the chamber, the iron particles were produced by evaporation in Ar gas at 80 Torr. The particles were collected on a glass plate 10 cm above the evaporation source. Immediately after the preparation of iron grains, the chamber was evacuated to 10⁻⁶ Torr, and SiO powder was evaporated from a W boat without exposure to air. The Fe particles covered with a SiO layer were removed from the vacuum system. The specimen was placed on the standard transmission electron microscope (TEM) grid and observed using an H-9000NAR TEM. The observation points were determined by placing the specimen on an electron microscopic holder, and the specimen was then heated in a furnace at 100°C in air. The alteration of the specimen was observed and analyzed at the same specimen position throughout the study.

3. Discussion

Figure 2(a) and (b) shows TEM images of Fe crystalline grains covered with a thin amorphous SiO layer and the specimen after heat-treatment at 100°C for 24 h in air, respectively. The same positions of the specimen are presented. As shown in the enlarged image in Fig. 2(a), the surface of each iron particle was covered with an amorphous SiO layer. The SiO was evaporated onto one side of the iron particles. However, the thickness difference seen in the enlarged image suggests that the entire iron grain surface was covered with a thin SiO layer. This suggests that the

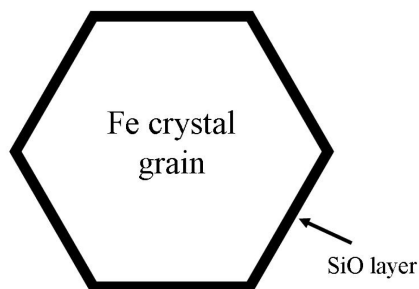


Fig. 1. Schematic model of Fe crystal grain covered with thin amorphous SiO layer.

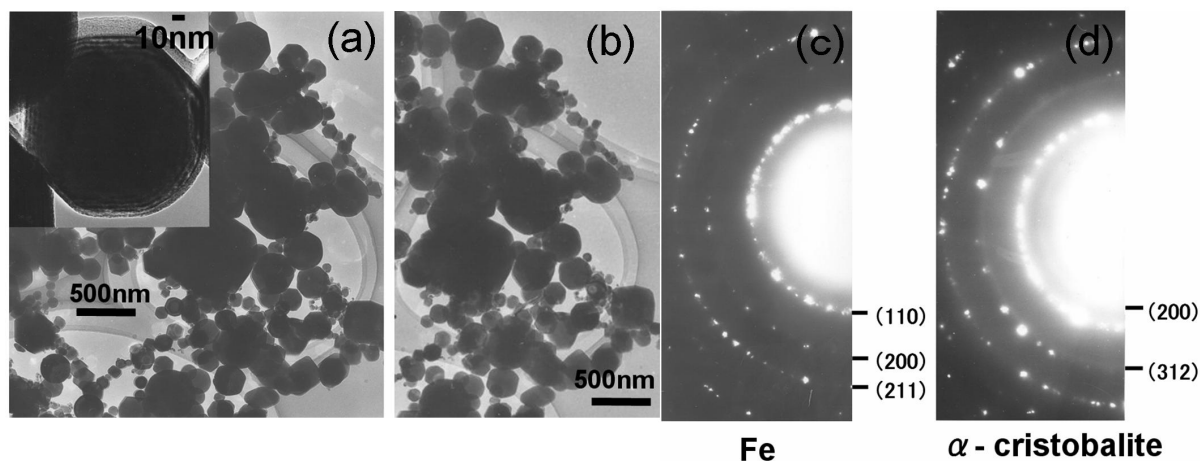


Fig. 2. Iron grain surface covered with SiO layer without exposure to air (a). The corresponding ED pattern (c) shows the iron crystal grain. The formation of a natural oxide layer was not observed. By heating the specimen at 100°C (b), the diffused cristobalite crystal rings appeared (d).

evaporated SiO covered the iron particle by the surface migration of SiO molecules. The ED patterns corresponding to the TEM images are shown in Fig. 2(c) and (d). The ED pattern in Fig. 2(c) shows both the existence of iron grains and the absence of oxide diffraction rings (Fe_3O_4), which generally appeared when a fresh iron particle was exposed to air (Tamura *et al.*, 2003). Therefore, it can be concluded that iron particle surfaces were covered with an amorphous layer of SiO. After heating for 24 h in air at 100°C, the ED pattern shows the existence of α -cristobalite crystallites, as can be seen in the diffraction pattern in Fig. 2(d). This suggests that the amorphous SiO layer was altered to a β -SiO₂ film.

Figure 3 shows an enlarged image of Fe grains after heating at 100°C for 24 h. The most common form of grown iron particles is a rhombic dodecahedron covered with 12 (110) planes (Kaito and Fujita, 1986). Such particles can be seen as different shapes along the zone axis. An enlargement of part of the image in Fig. 3(a) is shown in Fig. 3(b) where the (110) lattice of an iron particle can be clearly seen. The parallel fringes on the SiO₂ film region indicate the parallel growth of α -SiO₂ (202) and (101) planes, i.e., (110)_{Fe} // (101)_{SiO₂}. This suggests that the SiO layer was altered to a crystalline SiO₂ layer during heating at 100°C in air by the oxidation of Si crystallites on the surface of crystalline iron grains. The agreement between the lat-

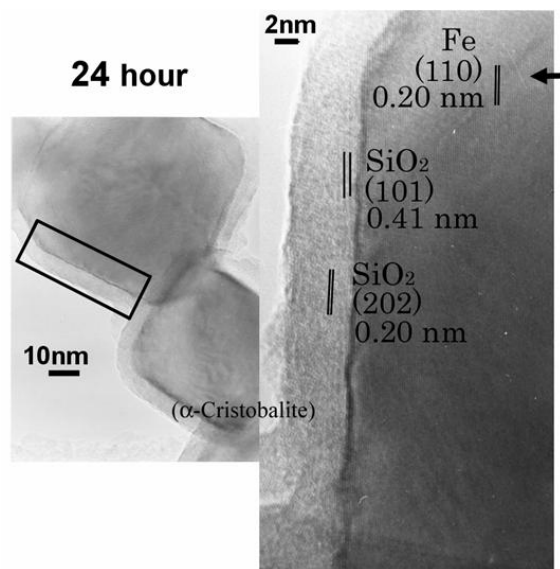


Fig. 3. After heating at 100°C for 24 h, the HRTEM image shows that (202) ($a_0 = 0.2019$ nm) and (101) (0.405 nm) crystals appeared parallel to the (110) surface of Fe. An arrow indicates the (110) lattice.

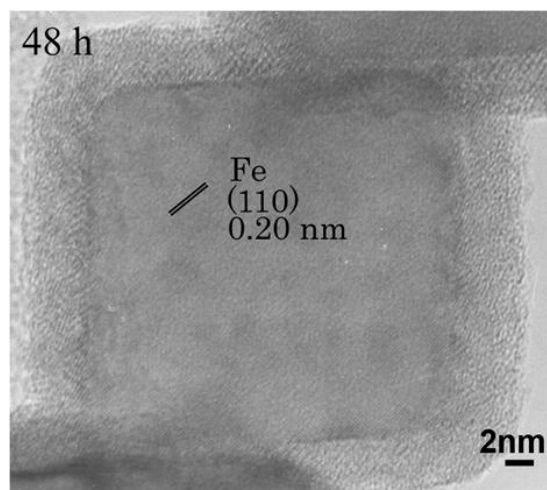


Fig. 4. The Fe particle was cubic because the (110) lattice can be seen at its corner. The growth of (101) and (202) cristobalite crystallites took place along the (110) plane of Fe.

tice of Fe(110) ($d = 0.202$ nm) and that of α -SiO₂ (202) ($d = 0.2019$ nm) is very good, i.e., the growth of the (202) layer took place at the least energy at the interface. The (101) lattice fringes were observed at part of the interface between the iron particle and the silicate layer.

Figure 4 also shows the cubic shape of a Fe grain covered with a SiO layer after heating at 100°C for 48 h. Since the (110) lattice of Fe can be seen along the edge of the cube, the iron particle shown in Fig. 4 was cubic and composed of six {100} planes. The (101) and (202) planes of SiO₂ remained parallel with the (110) plane of the Fe particle.

Since the mean free path of phonons in the crystal is of micrometer order, and the mean free path in the amorphous SiO layer is a few nanometers, i.e., the crystallite size, the oxidation energy of silicon crystallites in the SiO layer diffuses to the iron crystal grain. Therefore, the thermal energy due to the oxidation accumulates at the iron grain interface. This energy causes the crystallization of the SiO layer to SiO₂ due to the oxidation energy of the Si crystallites and diffusion reaction with air. The parallel relation at the interface between Fe and SiO₂ occurred due to a minimum in interface energy. This result supports the theory of low-temperature crystallization by Yamamoto *et al.* (2010). The oxidation energy of Si crystallites drives the crystallization of the SiO₂ layer.

Simple silicates evaporate to form SiO, the Mg atom, the Fe atom, and O₂ vapor and solid SiO, MgO, and FeO (Nuth *et al.*, 1999). Iron exists largely as metallic species in the mixture film of Fe and SiO (Suzuki *et al.*, 2000). If the iron grain is exposed in air, the grain surface is covered with magnetite phase and oxidation proceeds by the migration of iron atoms (Kaito *et al.*, 1973). The reaction between Fe and SiO takes place above 1100 K (Hallenbeck *et al.*,

1998). The crystalline SiO₂ growth on the iron grain surface in O₂ atmosphere results in low-temperature crystallization during the comet-formation process or cometary silicate crystallization as well as the previous experimental demonstration on the possibility of chemical-reaction-driven crystallization for the Mg-bearing silicate grain (Kaito *et al.*, 2007). The migration of iron atoms to the SiO layer hardly occurred at low temperature.

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C. Kaito (e-mail: kaito@se.ritsumei.ac.jp), A. Kumamoto, Y. Saito, and R. Ono