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Idealized numerical experiments on microscale eddies in the Venusian cloud layer

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Abstract

Three-dimensional microscale dynamics of convective adjustment and mixing in and around the Venusian lower cloud layer were investigated using an idealized Weather Research and Forecasting (WRF) model. As control parameters of the idealized experiment, the present work introduces an initial lapse rate in the convective layer and thermal flux associated with the infrared flux gap at cloud base. Eddy heat, material, and momentum fluxes increase in the convective layer with the increase of these two parameters. In the case of convective adjustment over a very short period, prior to formation of a large-scale convective cell, transient microscale eddies efficiently and rapidly eliminate the convective instability. In the case of convective mixing induced by cloud-based thermal flux, microscale eddies are induced around a thin unstable layer at the cloud base, and spread to the middle and upper parts of the neutral layer. For atmospheric static stability around 55 km, two types of fine structure are found: a wave-like profile induced by weak microscale eddies, and a profile locally enhanced by strong eddies.

Keywords: Venus; Atmospheric dynamics; Convection; Static stability

Background

On Venus, neutral and unstable layers were observed at height ranges of 50 to 55 km and less than 30 km (Seiff et al. 1980). The Vega 2 mission also observed zero and negative static stability in the lower clouds and near the surface (Young et al. 1987). The Venus Express radio science experiment (Tellmann et al. 2009) showed that the neutral and unstable layers around 50-km height extended to 45 km at high latitudes. Convective and turbulent motions have been observed in the Venusian neutral and unstable layers. The Vega balloon floated near 54 km (Sagdeev et al. 1986a), detecting vertical wind speeds of less than $1 \text{ m} \cdot \text{s}^{-1}$ during a quiet period, and greater than $3 \text{ m} \cdot \text{s}^{-1}$ during an active one (Sagdeev et al. 1986b). The observed convective heat flux ranged from 0 to $360 \text{ W} \cdot \text{m}^{-2}$ (Ingersoll et al. 1987; Crisp et al. 1990).

Micro- and mesoscale atmospheric dynamics, which are divided by a spatial scale of around 2 km in the case of the Earth (Orlanski 1975), are important topics in meteorology. Many works have conducted mesoscale simulations of the Venusian atmosphere. Baker et al. (1998,

2000a,b) simulated mesoscale convection and downward penetration with horizontal scales of 10 km around the 45-km level. However, microscale dynamics of turbulent convection on Venus have yet to be examined fully using a three-dimensional (3D) compressible and nonhydrostatic model.

The Weather Research and Forecast (WRF) model (Skamarock and Klemp 2008) has been used extensively to forecast weather and regional climate on Earth, and has also been applied to other planets (Lee et al. 2006; Richardson et al. 2007; Newman et al. 2011). Moeng et al. (2007) applied WRF to large-eddy simulations of the Earth's planetary boundary layer. Recently, Yamamoto (2011) applied WRF to 3D idealized microscale simulations of the Venusian atmosphere, examining transport processes of convective adjustment and mixing near the surface. The present work focuses on small eddies with scales of a few kilometers and durations of hours in the neutral and unstable layers of Venusian clouds (50 to 55 km), excluding meso- and global-scale convection in a small model domain. The goals are to clarify transport processes of microscale convective adjustment and mixing in the lower cloud layer, and to compare the microscale features with lander and balloon observations. Model assumptions and a description are given in section

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‘Methods’. Results are discussed in section ‘Results and discussion’ and summarized in section ‘Conclusions’.

Methods

Model assumptions

The atmospheric rotation period (6 to 8 days), radiative timescale (1 to 10 days; Crisp and Titov 1997; Titov et al. 2013), and Venusian solar day (117 days) are longer than that of the microscale eddy timescale (a few hours). Atmospheric radiative and rotational processes are not included, because the heating rate of approximately $1 \text{ K} \cdot \text{day}^{-1}$ (approximately 0.04 K h^{-1}) and the rotation rate are not significant for dynamical phenomena with timescales of a few hours. In the present idealized simulation, the initial lapse rate of potential temperature Γ_{LAP} and the turbulent thermal flux Q_B are set as tunable parameters controlling microscale dynamics in the neutral layer.

Positive Γ_{LAP} corresponds to the initial super adiabatic intensity leading to convective adjustment. The present work investigates turbulent eddies with temporal and spatial scales of a few hours and kilometers, resulting from convective adjustment (case A) under the initial unstable condition of positive Γ_{LAP} in the Venusian lower cloud.

According to Lee and Richardson (2011), as vertical resolution was increased in their radiative-convective model, the gradient of net upward infrared (IR) flux increased at the cloud base (dashed line in Figure 1). This implies that the IR-flux gap (thick gray arrow in the figure) is located at the cloud base and that heating has a locally sharp peak (ideally a delta function). If the IR-flux has a broad gradient, the heating rate is trivial for timescales of a few hours. In contrast, if the heating

from the IR-flux gap is a delta function at cloud base, the local heating rate is very large. Such localized heating there (where the local IR flux changes discontinuously) is not resolved by the present model. Instead of using the delta function, therefore, it is parameterized as the turbulent thermal flux at the bottom of the neutral layer. This flux may result from subgrid-scale eddies, *via* rapid changes of cloud morphology and heating with finer-scale condensation/evaporation. Although it is difficult to observe and estimate the turbulent flux at the boundary of the convective condensation and stable evaporation regions, this does not necessarily mean that subgrid-scale thermal flux Q_B , associated with a disturbed cloud base and IR gap, can be neglected.

The turbulent thermal flux Q_B in the vertical integration of the heat budget equation in the neutral layer is

$$Q_B \equiv [\theta'w']_{\text{Bottom}} \approx \frac{1}{\rho_{\text{Bottom}}} \int_{\text{Bottom}}^{\text{Top}} \rho \left\{ \frac{\partial[\theta]}{\partial t} + [\mathbf{u} \nabla[\theta] - [Q_{\text{RAD}}]] \right\} dz, \quad (1)$$

where $[\mathbf{u} \nabla[\theta]]$ is heat advection, $[Q_{\text{RAD}}]$ is radiative heating, and $\frac{\partial[\theta]}{\partial t}$ is the temperature tendency. Here, the square brackets indicate the area and time average over the model domain, and subscripts ‘Bottom’ and ‘Top’ indicate the bottom and top of the neutral layer, respectively. Equation 1 is derived by vertical integration of

$$\rho \left\{ \frac{\partial[\theta]}{\partial t} + [\mathbf{u} \nabla[\theta] - [Q_{\text{RAD}}]] \right\} = -\frac{\partial}{\partial z} \rho [\theta'w']. \quad (2)$$

The vertical integration of Equation 2 becomes

$$\int_{\text{Bottom}}^{\text{Top}} \rho \left\{ \frac{\partial[\theta]}{\partial t} + [\mathbf{u} \nabla[\theta] - [Q_{\text{RAD}}]] \right\} dz = -\rho_{\text{Top}} [\theta'w']_{\text{Top}} + \rho_{\text{Bottom}} [\theta'w']_{\text{Bottom}}. \quad (3)$$

Under the condition of $[\theta'w']_{\text{Top}} = 0$, Equation 3 becomes the vertically integrated budget equation (1).

The dynamical effect of long-lasting, large-scale motions is introduced as the advection term of large-scale motions in Equation 1. In the present study, ‘large scale’ means a horizontal size greater than that of the model domain (5 km). For an average over the large domain and a long period (square brackets in Equations 1 to 3), radiative heating should roughly balance the heat advection of large-scale motions. However, if there is transient subgrid-scale turbulence and an IR flux gap at the cloud base, $[\theta'w']_{\text{Bottom}}$ might be locally and transiently non-zero for a small domain. Thus, Q_B resulting from the subgrid-scale turbulence and IR flux gap should be considered as a thermal forcing parameter at cloud base in microscale simulations with a small domain. Given the assumptions that (1) $[Q_{\text{RAD}}]$ (approximately $0.04 \text{ K} \cdot \text{h}^{-1}$) induces large-scale circulation and convection $[\mathbf{u}]$ and

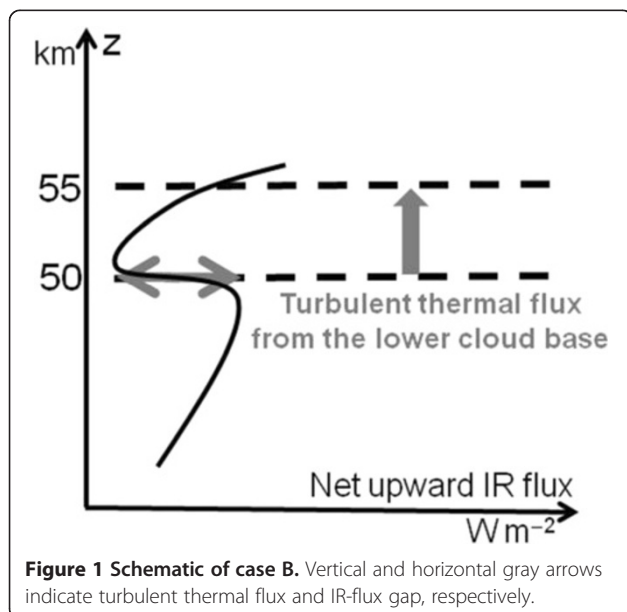


Table 1 Model conditions

Case	Γ_{LAP} ($\text{K} \cdot \text{km}^{-1}$)	Q_B ($\text{K m} \cdot \text{s}^{-1}$)
A02	0.2	0.000
A05	0.5	0.000
A10	1.0	0.000
A20	2.0	0.000
A30	3.0	0.000
B1	0.0	0.001 ($=1.44 \text{ W} \cdot \text{m}^{-2}$)
B2	0.0	0.004 ($=5.76$)
B3	0.0	0.016 ($=23.1$)
B4	0.0	0.064 ($=92.2$)
B5	0.0	0.256 ($=369$)

roughly balances their large-scale heat advection, (2) $[\theta' w']_{\text{Top}}$ is nearly zero at the neutral layer top, and (3) $[\theta' w']_{\text{Bottom}}$ is induced by transient subgrid-scale turbulence and the IR flux gap, Q_B is set as a tunable parameter in accord with the radiative simulation and balloon experiment results. The present work investigates turbulent microscale eddies associated with Q_B (case B).

Model description

Fully compressible and nonhydrostatic idealized simulations were conducted using the WRF Advanced Research model (WRF-ARW ver. 3.2), with the Arakawa C-grid for grid staggering in the Cartesian coordinate system. We set the 3D model domain of area 5×5 km with 50×50 grid points, and a height range of 50 to 58 km with 80 levels. Reference pressure and temperature at the 50-km level are 1,000 hPa and 350 K, respectively.

The acceleration owing to gravity g is $8.87 \text{ m} \cdot \text{s}^{-2}$, the gas constant R is $191.4 \text{ J} \cdot \text{kg}^{-1} \cdot \text{K}^{-1}$, the specific heat at constant pressure C_p is $904.0 \text{ J} \cdot \text{kg}^{-1} \cdot \text{K}^{-1}$, and the molecular weight of dry air is $44 \text{ g} \cdot \text{mole}^{-1}$. The rotational effect is not considered (the Coriolis parameter f is set to zero), because we focus on phenomena with time scales shorter than the atmospheric rotation. The third-order Runge-Kutta scheme is used for the time integration. Rayleigh damping for vertical flow with an inverse time scale of 0.2 s^{-1} and depth of 1,000 m from the model top is preset.

To compute subgrid-scale eddy diffusion for turbulent mixing, 1.5-order turbulent kinetic energy (TKE) closure is used (Section 4.2.3 of Skamarock et al. 2008). The diffusion coefficient is obtained from an empirical constant C_K , length scale l , and turbulent kinetic energy e . C_K is set to 0.1 (Xue et al. 2000), and l is defined by grid size, static stability, and turbulent kinetic energy. The subgrid-scale eddy diffusion could be sensitive to the empirical constants and grid scale, thereby influencing model results. At the present stage, it is difficult to apply higher resolution to the parameter sweep experiment under various initial and bottom-boundary conditions, because of the large computational resources required. Experiments investigating the sensitivity to grid size and subgrid-scale parameterization are necessary for future model validation.

The slip bottom boundary condition (i.e., drag coefficient set to zero) is used in all simulations, so the bottom momentum flux is set to zero. Fluxes of a passive tracer and heat at the bottom boundary are taken as zero and parameter Q_B ($\text{K m} \cdot \text{s}^{-1}$), respectively. Thus, the surface layer scheme is not applied in the model.

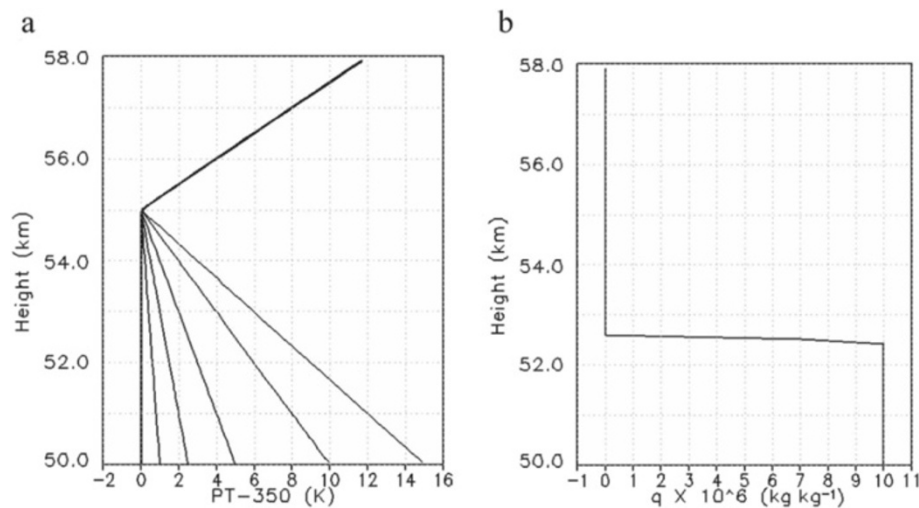
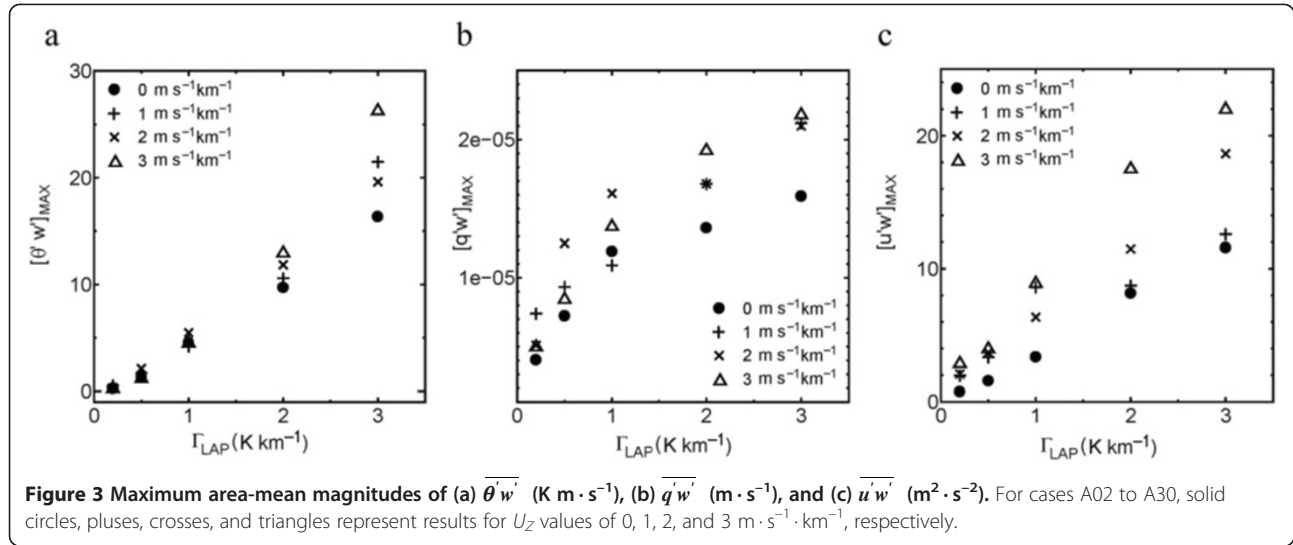


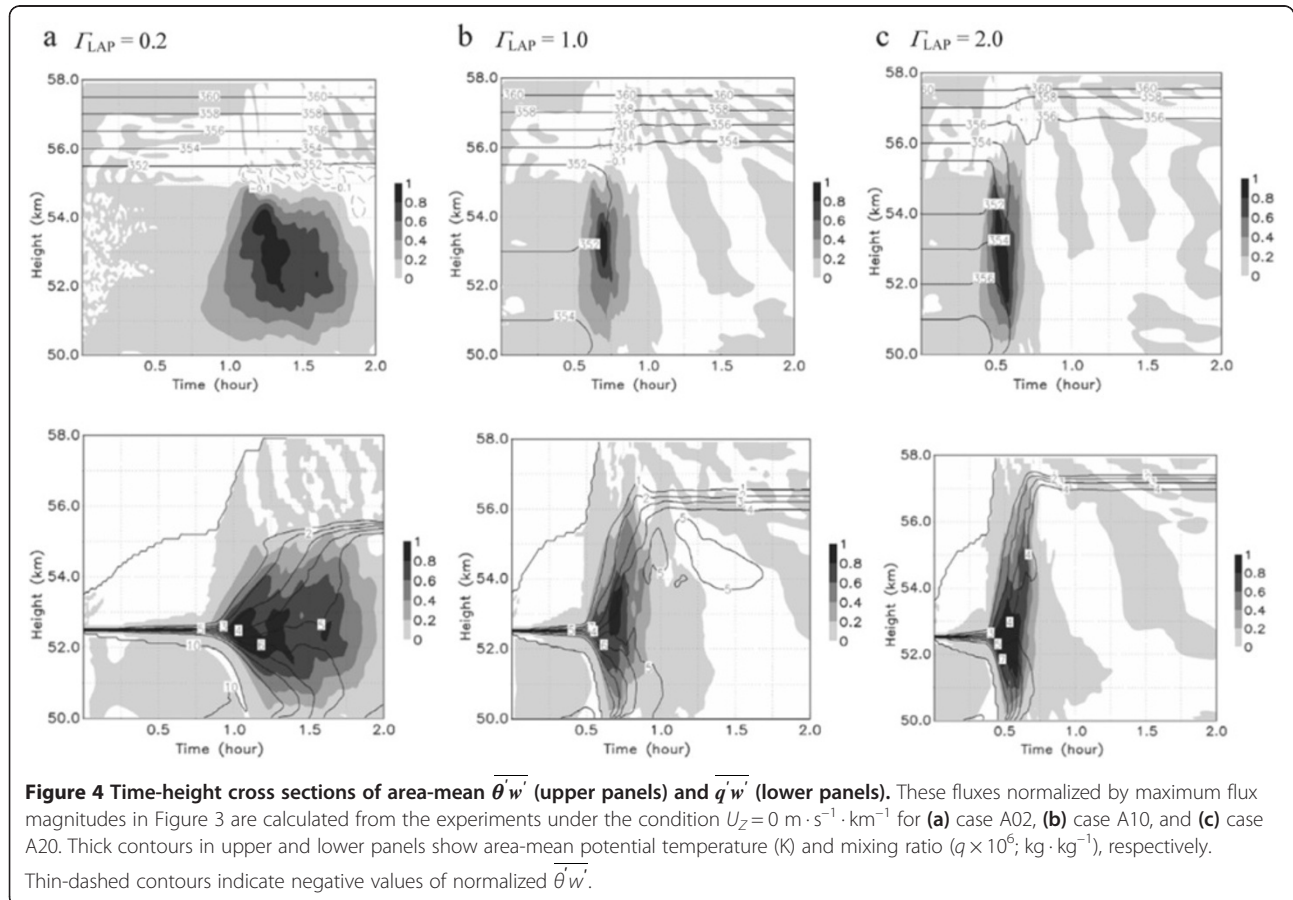
Figure 2 Initial vertical profiles of (a) potential temperature (K) and (b) passive tracer ($\text{kg} \cdot \text{kg}^{-1}$). The potential temperature difference from 350 K at 51 km is 0 K for cases B1 to B5, 0.8 K for case A02, 2 K for case A05, 4 K for case A10, 8 K for case A20, and 12 K for case A30.



A double periodic boundary condition is used along x (along the latitude line) and y (along the longitude line).

Γ_{LAP} is an initial lapse rate of potential temperature, and is defined as $-d\bar{\theta}/dz$ in an initially neutral or unstable

layer (50 to 55 km), which is capped by a stable layer (55 to 58 km) with buoyancy frequency $N \equiv [gd \ln \bar{\theta}/dz]^{1/2}$ of $0.01 \cdot s^{-1}$. In addition, a random perturbation is imposed on the mean temperature field at the lowest four grid levels, to initiate turbulent motion.



The experimental conditions of Γ_{LAP} and Q_B are listed in Table 1. Initial profiles of potential temperature θ and passive tracer mixing ratio q are shown in Figure 2. The convective adjustments are simulated in case A (A02 to A30, section ‘Convective adjustment in unstable layer (case A)’). Γ_{LAP} ranges from 0.2 to 3.0 $\text{K} \cdot \text{km}^{-1}$ in the neutral layer, and Q_B is set to zero. Microscale eddies induced by Q_B are simulated in case B (B1 to B5, section ‘Microscale eddies induced by turbulent thermal flux (case B)’). Q_B is varied from 0.001 to 0.256 $\text{K} \cdot \text{m} \cdot \text{s}^{-1}$. The wide range of Q_B , from 1.44 to 369 $\text{W} \cdot \text{m}^{-2}$ ($\rho = 1.594 \text{ kg} \cdot \text{m}^{-3}$, $C_p = 904 \cdot \text{kg}^{-1} \cdot \text{K}^{-1}$ at 50 km, Seiff et al. 1985), covers the convective heat flux observed in the Vega balloon experiments (0 to 360 $\text{W} \cdot \text{m}^{-2}$; Crisp et al. 1990; Ingersoll et al. 1987) and the IR-flux gap at cloud base (20 $\text{W} \cdot \text{m}^{-2}$; Lee and Richardson 2011). Γ_{LAP} is set to 0 $\text{K} \cdot \text{km}^{-1}$.

The vertical shear of the initial horizontal wind U_Z is set to 0 $\text{m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ in cases A and B. In addition, sensitivity simulations for initial wind shears 1, 2, and 3 $\text{m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ were also run to investigate the influence of the differentials between zonal winds in the lower and upper cloud layers. To confirm the presence of microscale eddies in the large model domain, domain size (5 × 5 km) in sections ‘Convective adjustment in unstable layer (case A)’ and ‘Microscale eddies induced by

turbulent thermal flux (case B)’ is extended by a factor of four (to 20 × 20 km) in the large-domain experiment (section ‘Sensitivity of microscale eddies to model domain size’).

Results and discussion

Convective adjustment in unstable layer (case A)

The initial lapse rate Γ_{LAP} ($= -d\bar{\theta}/dz$) is set as a control parameter of convective adjustment in case A. Maximum magnitudes of area-mean $\bar{\theta}'w'$, $\bar{q}'w'$, and $\bar{u}'w'$ increase with Γ_{LAP} as shown in Figure 3. Here, the overbar and prime indicate the zonal average and deviation from average, respectively. Eddy heat, material, and momentum transport are more efficient for larger Γ_{LAP} . The maximum $\bar{\theta}'w'$ and $\bar{u}'w'$ are sensitive to vertical shear of the initial zonal wind U_Z for large Γ_{LAP} . In particular, $\bar{u}'w'$ is amplified by increasing U_Z , although there are a few exceptions. For material transport, there are large differences of maximum $\bar{q}'w'$ among the four U_Z cases for the entire range of Γ_{LAP} .

Figure 4 shows time-height cross sections of area-mean $\bar{\theta}'w'$ and $\bar{q}'w'$ in cases A02 to A20. The initial convective adjustment causes upward eddy fluxes of heat

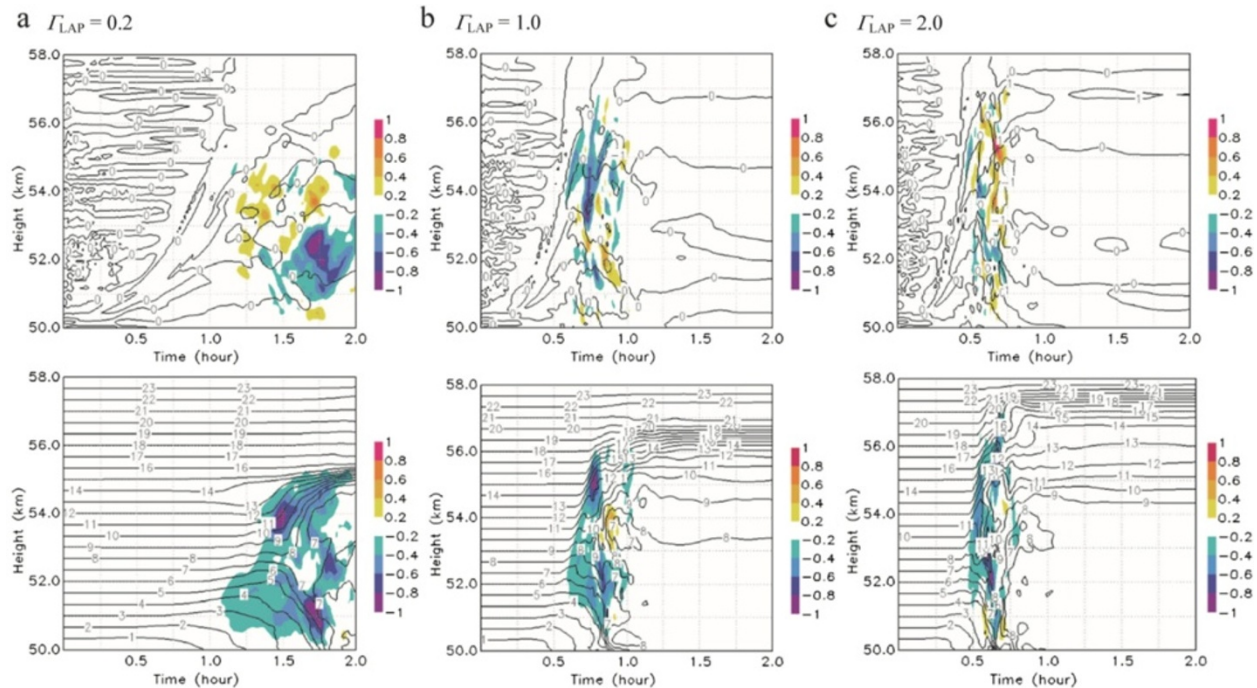


Figure 5 Time-height cross sections of area-mean $\bar{u}'w'$ in $U_Z = 0$ and 3 ($\text{m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$; upper and lower panels, respectively). These fluxes normalized by maximum flux magnitudes in Figure 3 are calculated from the experiments for cases (a) A02, (b) A10, and (c) A20. Contours show area-mean zonal wind velocity observed from the background-flow frame at 50 km ($\text{m} \cdot \text{s}^{-1}$).

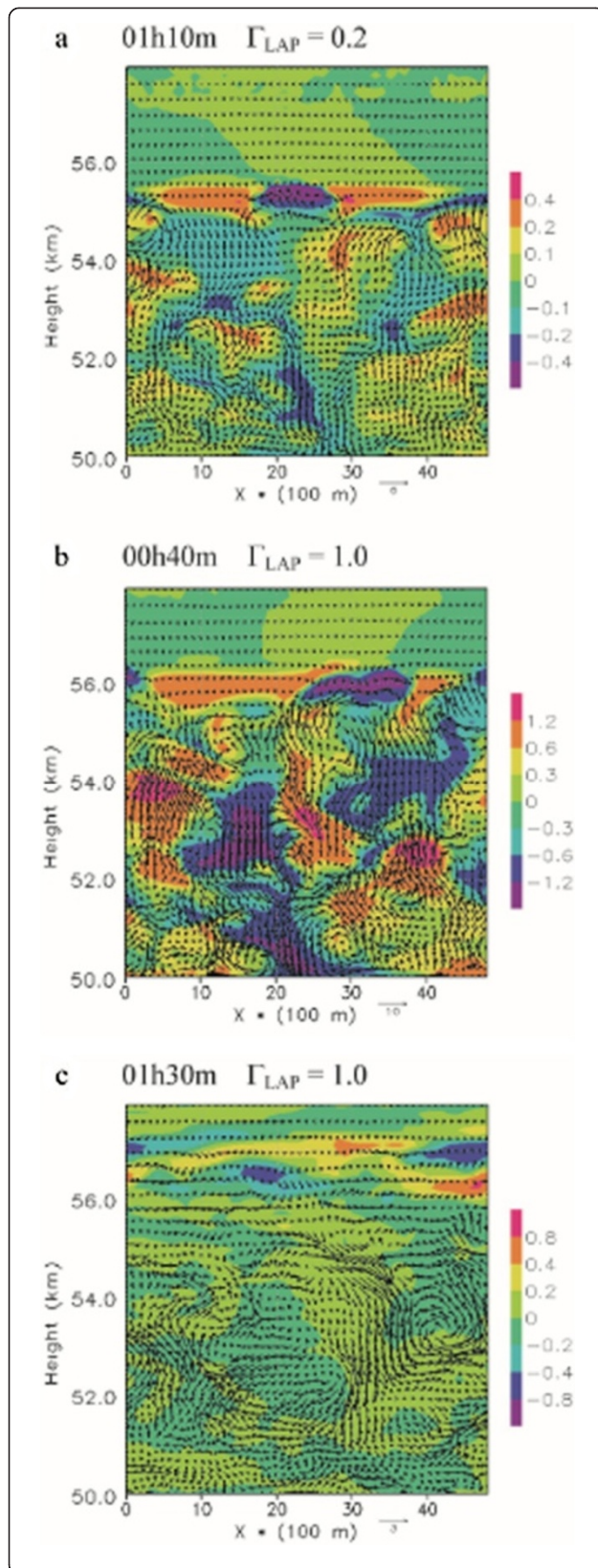


Figure 6 Zonal-vertical distributions of V and θ' at location $y = 1,000$ m. The eddy components of wind velocity V ($\text{m} \cdot \text{s}^{-1}$; arrows) and potential temperature θ (K; shaded) are calculated from the experiments under the condition $U_Z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ (a) at 70 min for case A02, (b) at 40 min for case A10, and (c) at 90 min for case A10.

and material. The convective adjustment occurs around 1 h in case A02, and 45 min in cases A10 and A20. With increasing Γ_{LAP} eddy kinetic energy increases in the neutral or unstable layer, and static stability ($d\theta/dz$) around 56 km weakens owing to penetration into the upper stable layer. Although the top of the mixed layer rises slightly with increasing Γ_{LAP} the convective adjustment does not significantly influence the upper stable layer throughout the simulation.

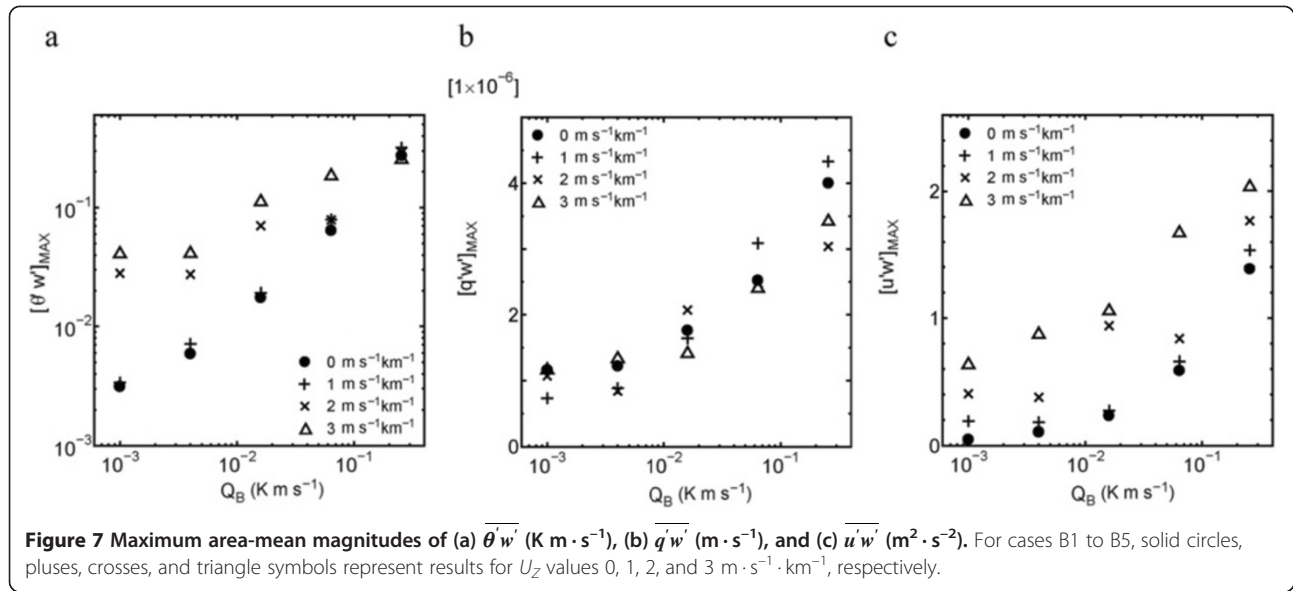
The eddy momentum flux is sensitive to the initial vertical shear of zonal flow U_Z . Figure 5 shows time-height cross sections of area-mean $\overline{u'w'}$ for U_Z values 0 and $3 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$. Large fluxes are evident at the time of convective adjustment. Structures of the mean flow and eddy momentum transport are complicated in the case of an initially zero wind-shear state ($U_Z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$), and velocities and fluxes fluctuate randomly with small amplitudes. In contrast, downward $\overline{u'w'}$ is predominant in the presence of a large initial wind shear ($U_Z = 3 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$). A narrow shear zone of the zonal flow develops around the mixed layer top (where contour intervals of mean zonal flow are dense), and weak shear persists in the upper part of the mixed layer after the convective adjustment.

Figure 6 shows snapshots of eddy wind velocity and potential temperature for weak and strong convective adjustments. Transient eddies with scales of a few kilometers are predominant in the unstable layer, but they are not found in the upper stable layer. After the strong convective adjustment penetrates the stable layer (to approximately 56 km), gravity waves are generated in its upper portion (Figure 6c), where the static stability has a wave-like profile (not shown).

Microscale eddies induced by turbulent thermal flux (case B)

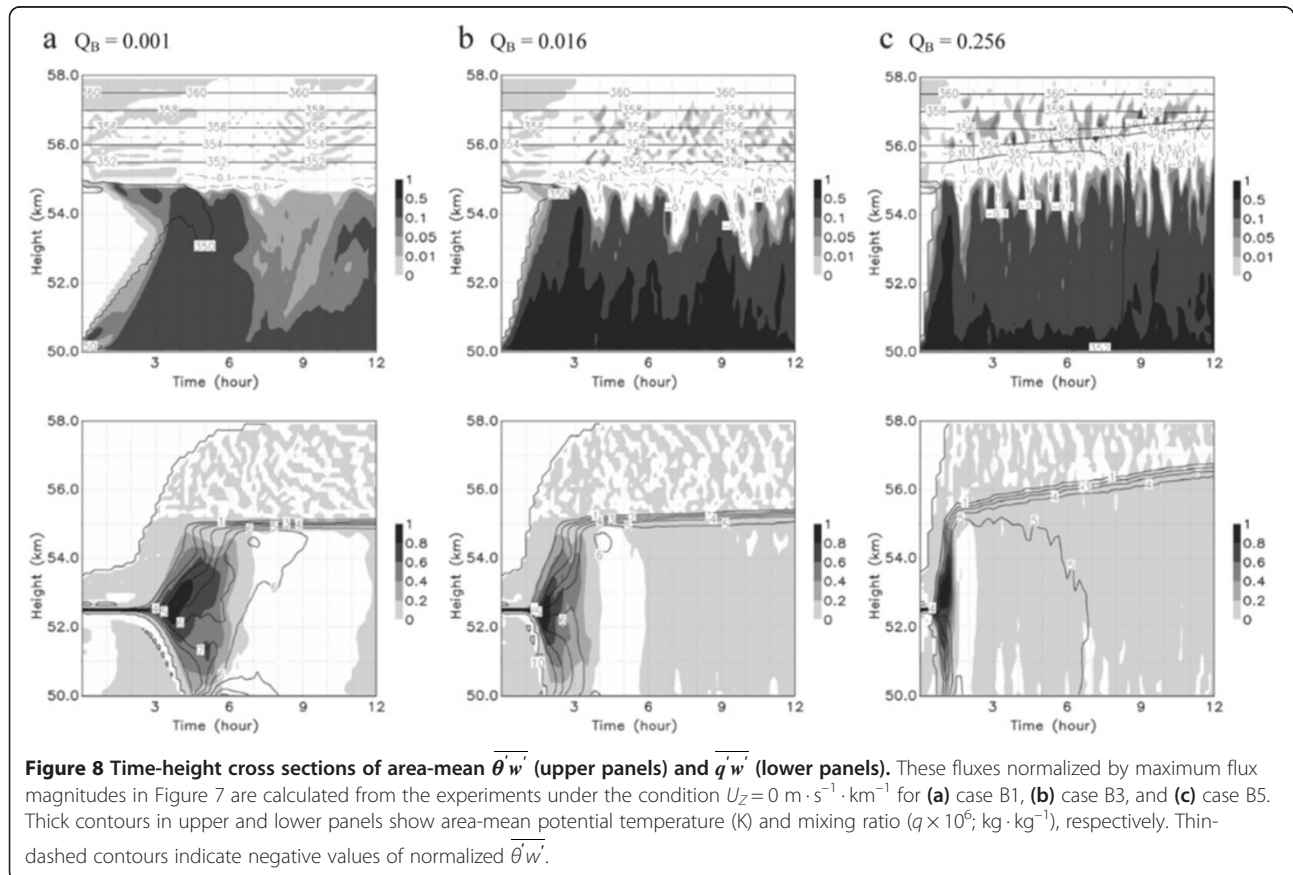
Q_B is set as a control parameter of convective mixing in case B. Figure 7 shows maximum magnitudes of area-mean $\overline{\theta'w'}$, $\overline{q'w'}$, and $\overline{u'w'}$. These flux magnitudes increase with Q_B . We see large differences of $\overline{\theta'w'}$ and $\overline{u'w'}$ among the four U_Z cases. Initially, strong shear of the zonal flow ($U_Z = 3 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$) is likely to produce large $\overline{\theta'w'}$ (left panel of Figure 7) and $\overline{u'w'}$ (right panel).

Time-height cross sections of area-mean $\overline{\theta'w'}$ for cases B1, B3, and B5 are shown in the upper panels of Figure 8.



In the case of small Q_B ($=0.001 \text{ K m s}^{-1}$), although the eddy upward heat flux is inefficient, weak mixing occurs throughout the simulation. For a Q_B of 0.016 K m s^{-1} , substantial upward heat transport is evident below 52 km and downward heat fluxes are apparent near 55 km

during the calculation. For large Q_B ($=0.256 \text{ K m s}^{-1}$), the vertical scale of the neutral layer grows gradually with time. The upward fluxes decrease rapidly with height and strong fluxes are confined near 50 km. Large Q_B at the bottom (values $\geq 0.016 \text{ K m s}^{-1}$) heats the



mixing layer *via* heat flux convergence of microscale eddies.

The passive tracer is transported rapidly upward with the initial convection (lower panels of Figure 8). After this, a uniform distribution forms in the neutral layer. A strong gradient of the tracer forms at the boundary between the neutral and stable layers.

Figure 9 shows time-height cross sections of area-mean $\overline{u'w'}$ for U_Z values of 0 and $3 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$. The area-mean eddy momentum flux and zonal flow change randomly in the neutral layer in the initially zero wind-shear state ($U_Z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ in the upper panels of Figure 9). In contrast, in cases of initially large shear (lower panels of the figure), downward $\overline{u'w'}$ is predominant upon the initial convection. Afterward, the shear zone of zonal flow develops at the boundary between the lower neutral and upper stable layers, and this zone persists because shear instability does not ensue. In contrast to small Q_B values ($\leq 0.016 \text{ K} \cdot \text{m} \cdot \text{s}^{-1}$), random disturbances of area-mean eddy momentum flux and zonal flow occurs in the mixing layer following the initial strongest momentum mixing with large Q_B ($=0.256 \text{ K} \cdot \text{m} \cdot \text{s}^{-1}$ in the lower right panel of Figure 9). Such strong wind shear forms *via* convective adjustment (Figure 5) and turbulent mixing (Figure 9).

Comparison with the balloon experiments is useful for discussing dynamical similarities with the observations and for qualitatively understanding the formation mechanism of fine structures; however, it is difficult to make a quantitative comparison in the idealized model configuration. Figure 10 shows the time series of vertical wind and eddy heat flux. The Vega balloon floated near 54 km (Sagdeev et al. 1986a) and detected vertical wind speeds less than $1 \text{ m} \cdot \text{s}^{-1}$ during a quiet period and greater than $3 \text{ m} \cdot \text{s}^{-1}$ during an active period (Sagdeev et al. 1986b). Downward winds were predominant as measured by the balloon, whereas there are both upward and downward motions in the model. This difference might be explained by the absence of large-scale circulation and radiation in the model, or by the dependence of vertical motions on the sampling locations. Observed convective heat flux ranged from 0 to $100 \text{ W} \cdot \text{m}^{-2}$ and increased abruptly to $360 \text{ W} \cdot \text{m}^{-2}$ (Ingersoll et al. 1987; Crisp et al. 1990). At 55 km ($\rho = 0.9207 \text{ kg} \cdot \text{m}^{-3}$, $C_p = 859 \text{ kg}^{-1} \cdot \text{K}^{-1}$; Seiff et al. 1985), the observed heat flux $100 \text{ W} \cdot \text{m}^{-2}$ equals that for $\theta'w' = 0.126 \text{ K} \cdot \text{m} \cdot \text{s}^{-1}$. The eddy heat fluxes fluctuate slightly during the simulation, but large fluxes appear intermittently. The maxima of 30, 320, and $480 \text{ W} \cdot \text{m}^{-2}$, which are larger than Q_B , occur in cases B3, B4, and B5, respectively. The abrupt increase of convective

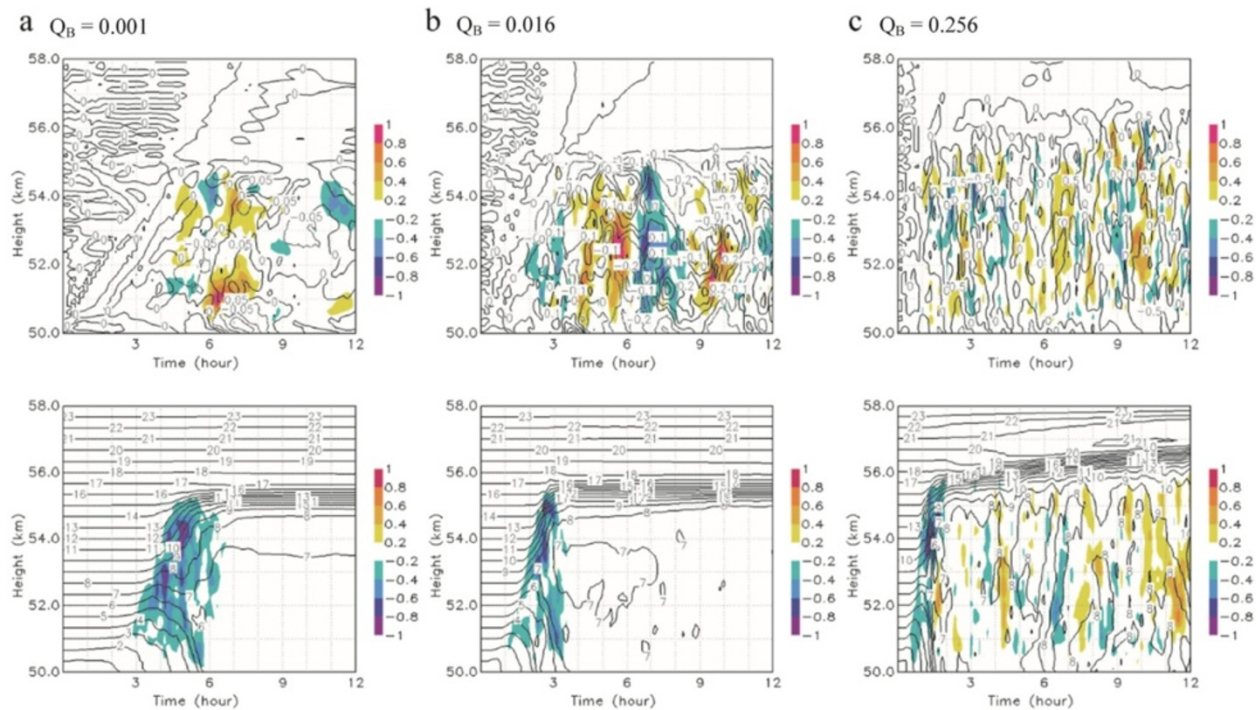


Figure 9 Time-height cross sections of area-mean $\overline{u'w'}$ in $U_Z = 0$ and $3 \text{ (m} \cdot \text{s}^{-1} \cdot \text{km}^{-1})$; upper and lower panels, respectively). These fluxes normalized by maximum flux magnitudes in Figure 7 are calculated from the experiments for cases (a) B1, (b) B3, and (c) B5. Contours show area-mean zonal wind velocity observed from the background-flow frame at 50 km ($\text{m} \cdot \text{s}^{-1}$).

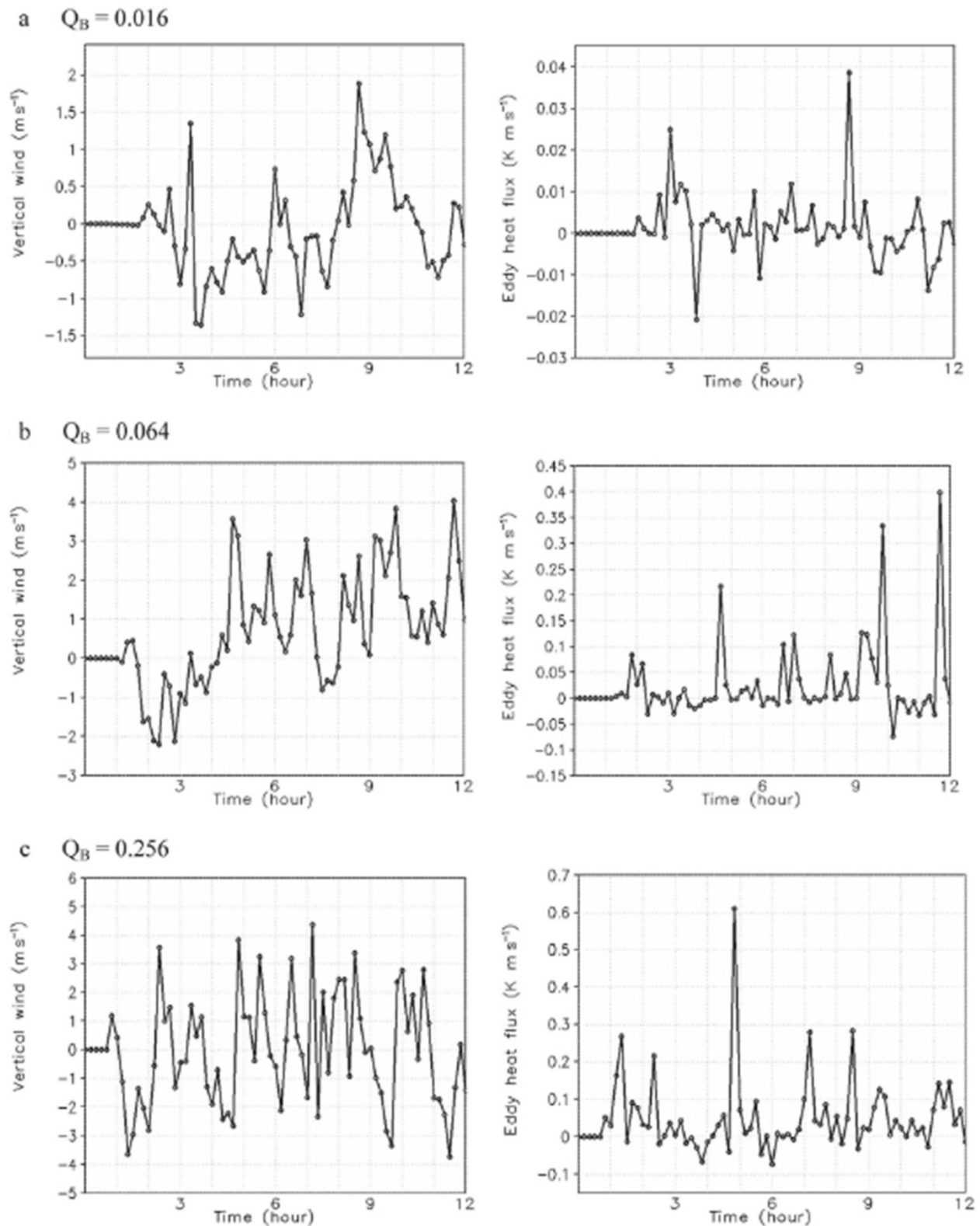


Figure 10 Time histories of vertical wind velocities ($\text{m} \cdot \text{s}^{-1}$; left panels) and eddy heat fluxes ($\text{K} \cdot \text{m} \cdot \text{s}^{-1}$; right panels). These are plots of output data at location $x = 1,000$ m and $y = 1,000$ m, and 54-km level under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ for (a) case B3, (b) case B4, and (c) case B5.

heat flux in case B4 is comparable to that observed in the Vega experiment ($360 \text{ W} \cdot \text{m}^{-2}$). This indicates the possibility that intermittent microscale eddies generate such an increase.

Figure 11 shows vertical profiles of the atmospheric static stability. Wave-like fluctuations are evident in the stable layer above 55 km in the upper panels of the figure. These correspond to a temperature pattern induced by gravity waves of short vertical wavelength of approximately 1 km and small amplitude of approximately 0.1 K in the upper stable layer (gray ellipses in upper panels of Figure 12). These gravity waves are forced by turbulent microscale eddies and not by strong deep convective penetration of the upper stable layer. Such small-scale gravity waves were not generated by the large-scale convection cited in Baker et al. (1998).

However, local and frequent penetrations maintained by turbulent microscale eddies might force such gravity waves. In the lower panels of Figure 11, static stability increases locally to approximately $10 \text{ K} \cdot \text{km}^{-1}$ around 56 km. This locally enhanced structure is caused by a temperature patch resulting from the strong turbulent eddies (gray ellipses in lower panels of Figure 12). Such local increase of stability is also seen above 55 km in the Pioneer Venus and Vega observations (Young et al. 1987). The transient nature and local amplitude of microscale eddies around 55 km determines the static stability pattern. When the eddies strengthen (weaken) around the top of the neutral layer, the locally enhanced (wave-like) structure is likely to appear in the upper stable layer.

In contrast to the upper stable layer, transient eddies of scales 1 to 2 km are predominant in the convective

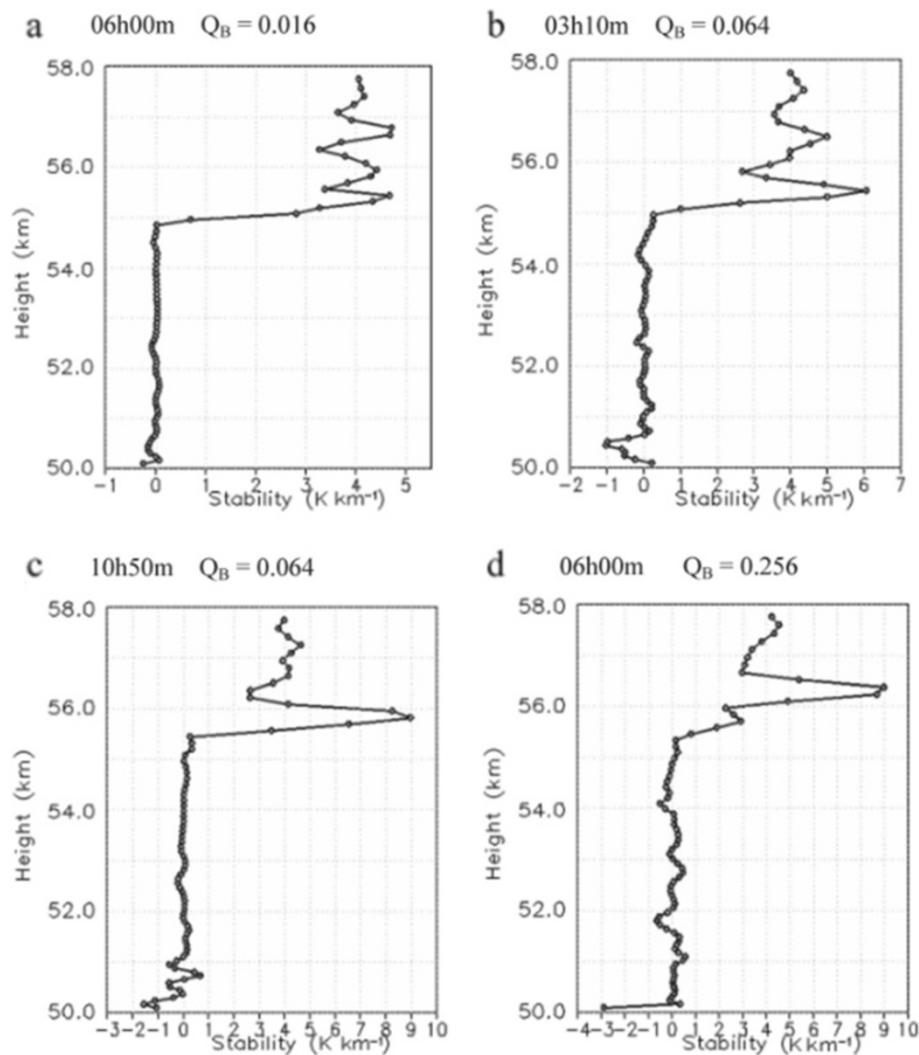


Figure 11 Snapshots of vertical profiles of static stabilities at location $x = 1,000 \text{ m}$ and $y = 1,000 \text{ m}$. The static stabilities are calculated from the experiments under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ at (a) 360 min for case B3, (b) 190 min for case B4, (c) 650 min for case B4, and (d) 360 min for case B5.

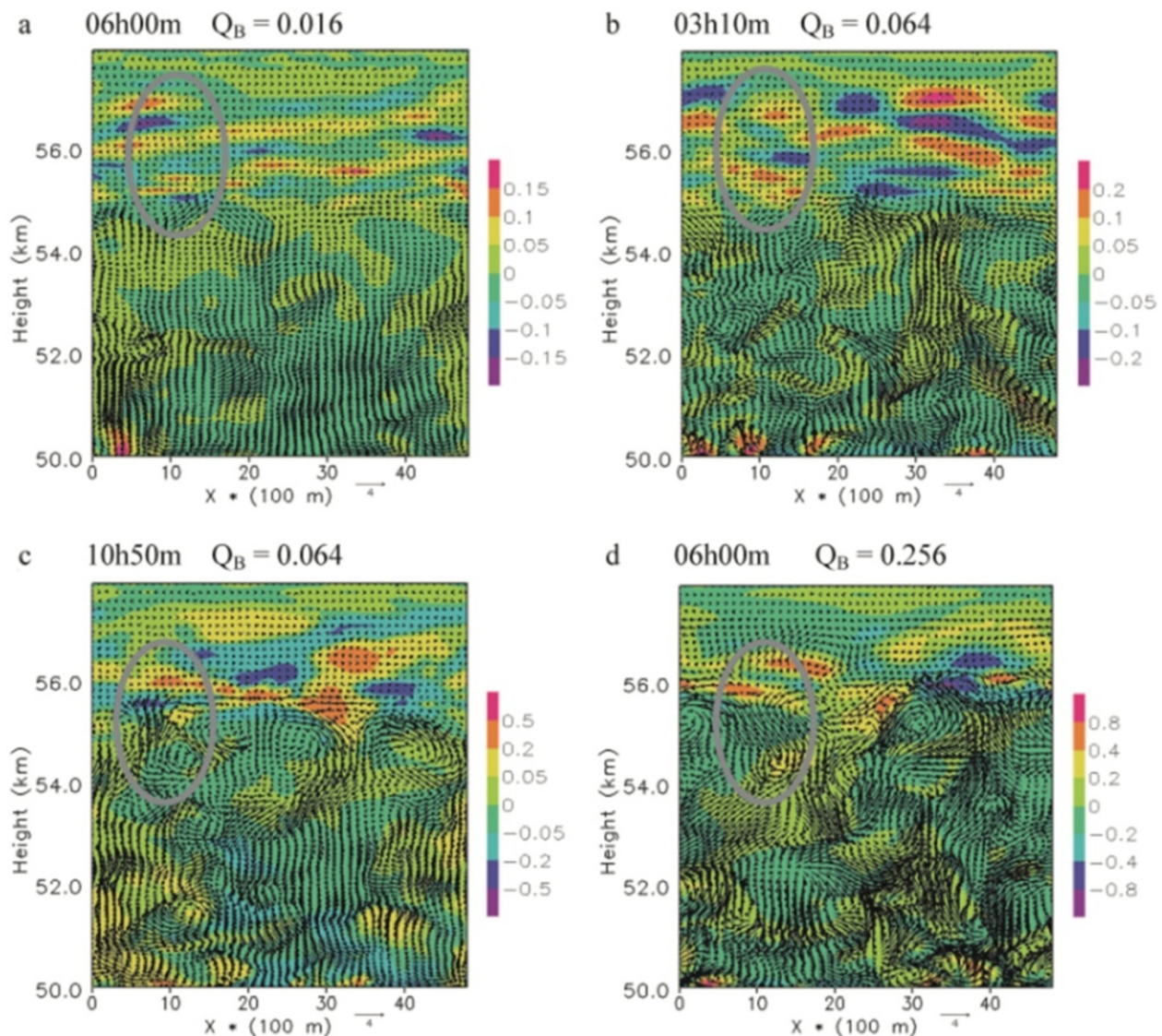


Figure 12 Zonal-vertical distributions of V' and θ' at location $y = 1,000$ m. The eddy components of wind velocity V' ($\text{m} \cdot \text{s}^{-1}$; arrows) and potential temperature θ' (K; shaded) are calculated from the experiments under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ at (a) 360 min for case B3, (b) 190 min for case B4, (c) 650 min for case B4, and (d) 360 min for case B5.

layer. However, eddy temperatures in the neutral layer are smaller than those above 56 km (Figure 12). Thus, the small temperature deviation associated with microscale convective motions results in small fluctuations of static stability in the convective layer (50 to 55 km; Figure 11).

Sensitivity of microscale eddies to model domain size

Sensitivity to model domain size is investigated for several cases to confirm the presence of the microscale eddies in a large domain of 20×20 km. The domain area is 16 times larger than in sections 'Convective adjustment in unstable layer (case A)' and 'Microscale eddies

induced by turbulent thermal flux (case B)'. Figure 13 shows time-height cross sections of area-mean $\overline{\theta' w'}$, $\overline{q' w'}$, and $\overline{u' w'}$ for large-domain experiments of cases A10 and B3 under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$. In the case of convective adjustment, time histories and magnitudes of eddy heat and material fluxes in the large domain experiment (case A10L) are nearly the same as those of case A10. The area-mean zonal wind becomes smooth in case A10L. The area-mean wind shear inducing the eddy momentum flux also weakens because the zonal flow is averaged over the large domain. Thus, the area-mean maximum eddy momentum flux for case A10L is one-third less than that of case A10. In the case of convective

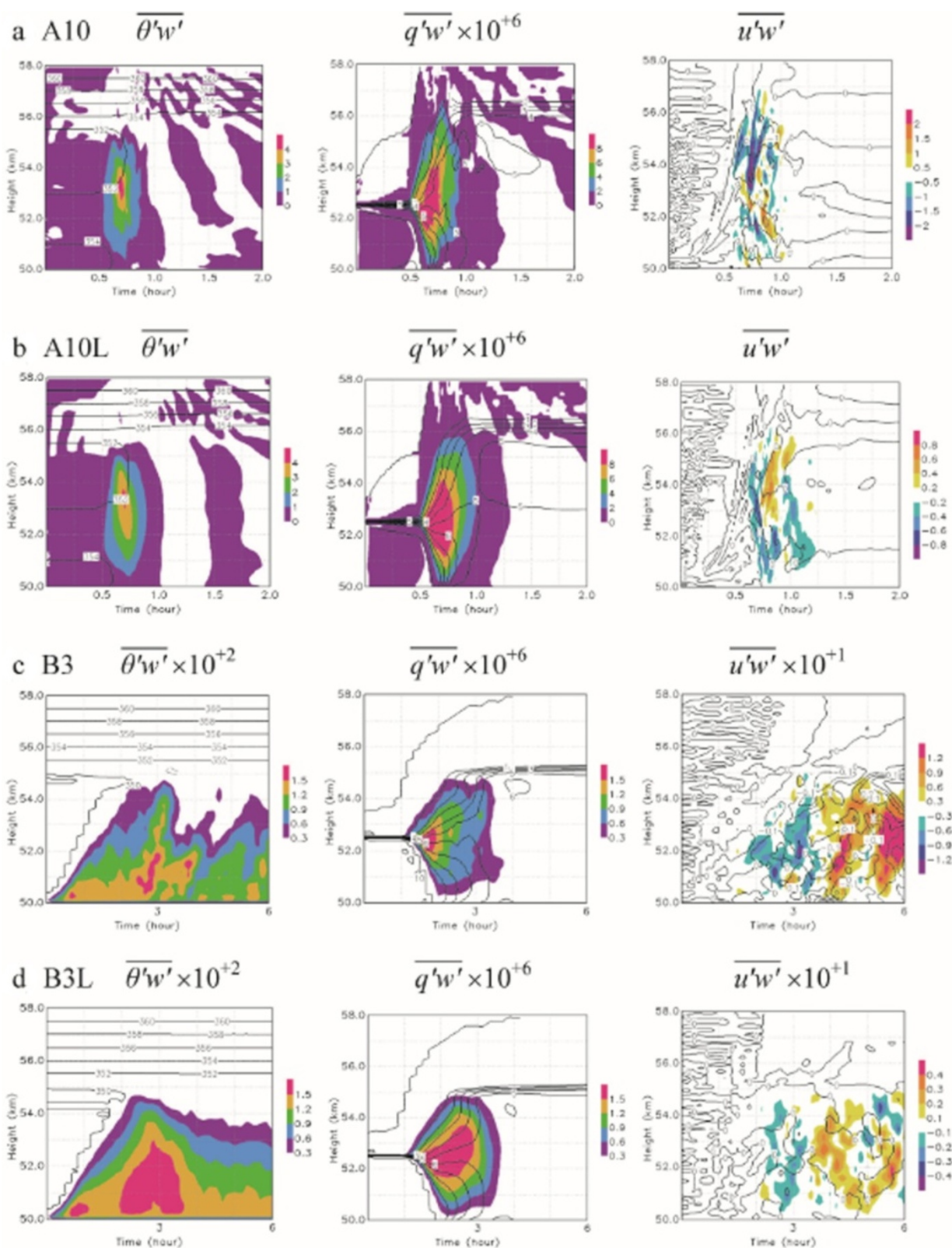


Figure 13 (See legend on next page.)

(See figure on previous page.)

Figure 13 Time-height cross sections of area-mean $\overline{\theta'w'}$ (left panels), $\overline{q'w'}$ (middle panels), and $\overline{u'w'}$ (right panels). These fluxes (shaded) are calculated from the experiments under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ for (a) case A10, (b) case A10L, (c) case B3, and (d) case B3L. Thick contours show area-mean potential temperature (K; left panels), mixing ratio ($q \times 10^6; \text{kg} \cdot \text{kg}^{-1}$; middle panels), and wind speed ($\text{m} \cdot \text{s}^{-1}$; right panels). 'L' appended to the case name indicates large-domain experiment.

mixing induced by Q_B , maximum eddy heat and material fluxes in case B3L are the same as those in B3 (differences were only a few percent). However, the large flux regions (red shading) are somewhat wider in the large-domain experiment (Figure 13d). The random fluctuation and vertical shear of the mean zonal wind becomes smooth, and the area-mean momentum flux associated with the mean wind shear also weakens in the large-domain experiment.

Figure 14a,b shows horizontal distributions at 54 km of eddy components of wind ($\text{m} \cdot \text{s}^{-1}$; arrows) and potential temperature (K; shaded) for convective adjustment in case A10L, and strong convective mixing in case B3L. Turbulent eddies with sizes equal to and smaller than 5

km are predominant in the large-domain experiments. Strong microscale eddies are relatively dense in A10L. In contrast, microscale eddies are relatively sparse and weak in B3L. Figure 14c,d shows amplitudes of the Fourier component of potential temperature for each zonal wavelength, averaged over an extended sampling area of $25.6 \times 25.6 \text{ km}$. Here, we assumed that the wave structures are cyclically repeated outside the model domain. This was done to analyze the output data with a double periodic boundary under the same condition in which the sampling number was 2^8 and the data grid interval was 100 m in the fast Fourier transform. Signals with zonal wavelengths of 2.5 and 5 km are predominant

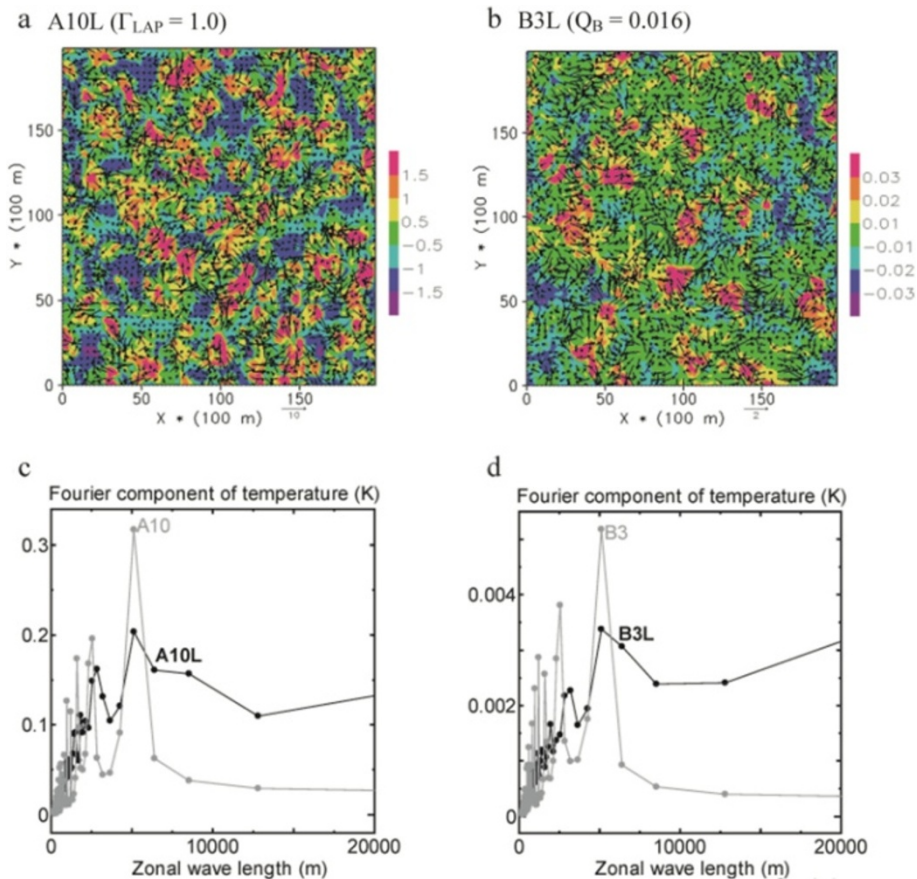


Figure 14 Horizontal distributions (a,b; upper panels) and amplitudes of Fourier components (c,d; lower panels) of potential temperature. These are plots of eddy components of wind ($\text{m} \cdot \text{s}^{-1}$; arrows) and potential temperature (K; shaded) at 54 km under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ at (a) 40 min for case A10L and (b) 180 min for case B3L. Black and gray marks in (c,d) show amplitudes of potential temperatures in cases A10L and B3L and in cases A10 and B3, respectively.

in these experiments. The strongest signals are enhanced by 50% in the small-domain experiments, because of wave structure repetition under the cyclic condition every 5 km. In contrast, strong small-scale eddies randomly (not cyclically) occur in the large domain. The strong but sparse small-scale eddies along the x direction are partly decomposed to Fourier components of large scales >5 km. Thus, the strongest signals at 5 km become small in the large-domain experiments, relative to the small-domain experiments.

When a sub-domain of $x=0$ to 5 km and $y=0$ to 5 km is considered in the large-domain experiments, signals of zonal wavelengths >5 km do not predominate in the sub-domain. Zonal-vertical cross sections of eddies in the sub-domain of 5-km size in the large-domain experiments (Figure 15) are similar to those of the small-domain experiments (Figure 12). Weak turbulent eddies in case B3L induce gravity waves in the upper stable layer, resulting in a wave-like profile of static stability above 55 km. Conversely, strong microscale eddies in case B4L produce temperature patches on scales of approximately 2 km, which are locally enhanced around 55.5 km at $x=1,000$ m. This produces a locally enhanced profile of static stability. Thus, these results in the large-domain simulation are consistent with those in section ‘Microscale eddies induced by turbulent thermal flux (case B)’.

As mentioned above, 3D microscale eddies are predominant in the large-domain simulations (20×20 km) as in the simulations with domains of 5×5 km. Model domain size does not greatly influence the turbulent mesoscale eddies and their heat and material transports, although the momentum flux magnitude decreases in the initially zero wind-shear experiments. In the presence of turbulent microscale eddies, the following two processes are important: (i) In the case of convective

adjustment over very short periods (10 to 30 min), before a large-scale convection cell forms, the transient turbulent eddies efficiently and rapidly eliminate the initially unstable state. (ii) In the case of convective mixing induced by Q_B , because forcing of Q_B is confined to the cloud base and is maintained, strong eddies of 1-km length are induced around a thin unstable layer at the cloud base (around 50 km in Figure 11) and spread to the middle and upper parts of the neutral layer. Thus, the microscale eddies are maintained, and eddy heat fluxes are large in the lower part of the neutral layer. Such microscale features in (i) and (ii) are not found in previous mesoscale simulations.

Conclusions

The present work investigated 3D microscale dynamics of convective adjustment and mixing in and around the Venusian lower cloud layer, and examined heat, material, and momentum transport processes of microscale eddies. In an idealized WRF model, initial lapse rate and bottom thermal flux are given as control parameters of microscale adjustment and mixing in the lower cloud layer between 50 and 55 km. Eddy heat, material, and momentum fluxes are enhanced in the lower cloud layer with increasing initial lapse rate in the unstable layer and bottom thermal flux in the neutral layer. When shear of the zonal wind is initially present, a strong thin shear zone of zonal wind is able to form around the top of the convective layer after the initial strong convective motions, and momentum flux magnitudes somewhat increase. The eddy heat and material fluxes are not sensitive to model domain size. However, because the random fluctuation and vertical shear of the mean zonal wind become smooth in the large-domain experiment, area-mean momentum flux associated with the shear weakens in the initially zero wind-shear experiments.

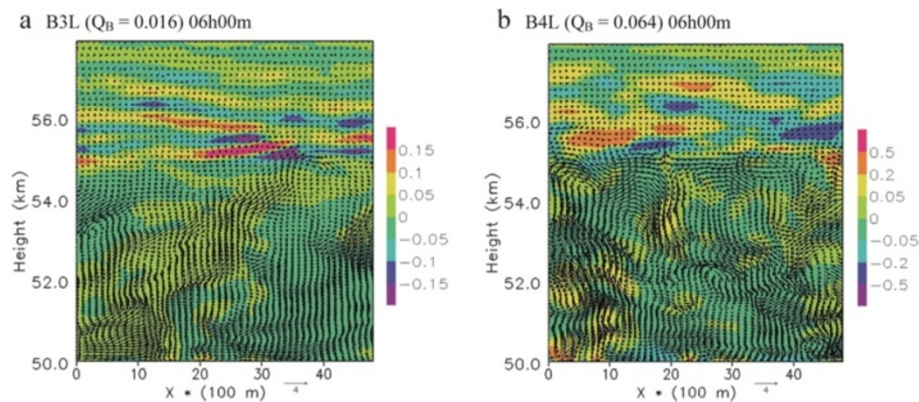


Figure 15 Zonal-vertical distributions of V' and θ' at location $y = 1,000$ m. The eddy components of wind velocity V ($\text{m} \cdot \text{s}^{-1}$; arrows) and potential temperature θ (K; shaded) at 360 min are calculated from the experiments under the condition $U_z = 0 \text{ m} \cdot \text{s}^{-1} \cdot \text{km}^{-1}$ for (a) case B3L and (b) case B4L.

Thus, we should carefully ascertain the dependence of vertical eddy momentum transport on domain size.

If convective adjustment and Q_B -induced mixing occur in the Venusian lower cloud, microscale eddies should be considered in eddy heat and material transport processes. In the case of convective adjustment over very short periods (10 to 30 min), before a large-scale convection cell forms, the transient microscale eddies efficiently and rapidly eliminate the initially unstable state. In the case of convective mixing induced by Q_B , because forcing of Q_B is confined to the cloud base and is maintained, strong small eddies are initially induced around a thin unstable layer at the cloud base and spread to the middle and upper parts of the neutral layer.

Weak turbulent eddies induce microscale temperature patches associated with gravity waves in the upper stable layer, and thereby form a wave-like profile of static stability above 55 km. Such small turbulent eddies may contribute to the forcing mechanism of gravity waves. Conversely, strong microscale eddies produce locally enhanced structures of atmospheric stability. Since amplitudes of the microscale turbulent eddies are large around the mixing layer top, the locally enhanced structures are likely to appear in the upper stable layer.

Competing interests

The author declares that he has no competing interests.

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