Spatial distribution of F-net moment tensors for the 2005 West Off Fukuoka Prefecture Earthquake determined by the extended method of the NIED F-net routine

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The 2005 West Off Fukuoka Prefecture Earthquake with a Japan Meteorological Agency (JMA) magnitude (M_{MA}) of 7.0 occurred on March 20, 2005. We determined moment tensor solutions, using a surface wave with an extended method of the NIED F-net routine processing. The horizontal distance to the station is rounded to the nearest interval of 1 km, and the variance reduction approach is applied to a focal depth from 2 km with an interval of 1 km. We obtain the moment tensors of 101 events with (M_{MA}) exceeding 3.0 and spatial distribution of these moment tensors. The focal mechanism of aftershocks is mainly of the strike-slip type. The alignment of the epicenters in the rupture zone of the main-shock is oriented between N110°E and N130°E, which is close to the strike of the main-shock's moment tensor solutions (N122°E). These moment tensor solutions of intermediate-sized aftershocks around the focal region represent basic and important information concerning earthquakes in investigating regional tectonic stress fields, source mechanisms and so on.

Key words: The 2005 West Off Fukuoka Prefecture Earthquake, moment tensor, broadband seismometer, seismicity, F-net.

1. Introduction

The 2005 West Off Fukuoka Prefecture Earthquake M_{JMA}7.0 that occurred at 1:53 a.m. (UT) on March 20, 2005 is the first M7-class intra-plate earthquake in the northern Kyushu district after F-net, the broadband seismograph network of the National Research Institute for Earth Science and Disaster Prevention (NIED) was established with a dense and homogeneous distribution all over Japan. The seismicity around the northern Kyushu district is very low. An M7-class earthquake occurred near Iki and Tsushima in 1700, and this is the only earthquake that occurred historically around this region (The Headquarters for Earthquake Research Promotion, 1997). Seno (1999) estimated the stress field around Japan. This model indicates the maximum compressive stress direction from E-W to ENE-WSW in the northern Kyushu district. It is consistent with the maximum compressive stress direction indicated by the focal mechanisms of background seismicity, which is determined by NIED moment tensor routine processing (Fig. 1).

Earthquake focal mechanisms represent basic and important information in the seismology and have been utilized to understand the regional tectonic stress fields and the source mechanisms of large earthquakes, and to simulate the strong motion and so on.

In order to estimate the distribution of the source fault in and around the focal area of the 2005 west off Fukuoka prefecture earthquake, Ito *et al.* (2006) mainly employed

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the centroid moment tensor inversion method to calculate the centroid location, time and moment tensor solutions of the main shock and aftershocks with a magnitude more than 4, velocity and acceleration seismograms of the F-net broadband seismometers and Hi-net tiltmeters. They also estimate the direction of rupture propagation of the main shock and finally discuss the initiation and termination of the rupture of the main shock controlled by fault bends.

In this study, we thoroughly investigate almost all the spatial distribution of moment tensors of the main-shock and their intermediate-sized aftershocks with a magnitude less than 4, through the extended method of the NIED F-net routine processing. In this analysis, moment tensor solutions and centroid depths were only calculated by using waveforms observed at three stations, with the fixed epicenters in the JMA Catalogue. We finally discuss the relationship between the distribution of moment tensor solutions and the coseismic-slip area of the main shock.

2. Moment Tensor Inversion

Waveform data are obtained from the F-net, the broadband seismograph network of NIED installed at intervals of 100 km in Japan (Okada *et al.*, 2004) (Fig. 1). A three-component broadband seismometer (STS-1/2) and a three-component strong motion velocity-type seismograph (VSE-355G2/G3) have been respectively installed at each station. We have been routinely analyzing earthquakes of magnitude greater than 3.5 based on the unified hypocenter catalog maintained by the Japan Meteorological Agency (JMA) (Fukuyama *et al.*, 1998). All results of the moment tensor analysis using surface wave are available or accessible

Depth	Thickness	Vp	Vs	Density	Qp	Qs
(km)	(km)	(km/s)	(km/s)	(kg/m^3)		
0	3	5.500	3.140	2300	600	300
3	15	6.000	3.550	2400	600	300
18	15	6.700	3.830	2800	600	300
33	67	7.800	4.460	3200	600	300
100	125	8.000	4.570	3300	600	300
225	100	8.400	4.800	3400	600	300
325	100	8.600	4.910	3500	600	300
425	60	9.300	5.310	3700	600	300
485	60	9.713	5.467	3850	600	300
545	60	9.914	5.624	3950	600	300
605	60	10.120	5.781	4030	600	300
665	60	10.880	6.041	4380	600	300
725		11.060	6.201	4500	600	300

Table 1. Structure model.

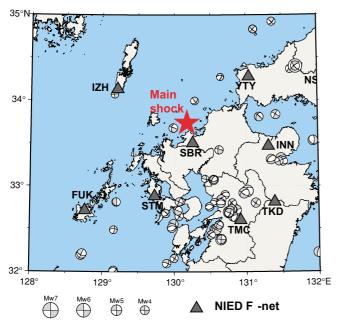


Fig. 1. Station distribution and seismicity around northern Kyushu, Japan. Focal mechanisms less than 25 km in depth occurred from January 1997 until March 20, 2005. The red star indicates the epicenter of the 2005 West Off Fukuoka Prefecture Earthquake, which occurred on March 20, 2005 (JMA, 2005). Black triangles show the NIED F-net stations.

through the Internet (http://www.fnet.bosai.go.jp/). These source parameters are used as the initial condition in the waveform inversion for the rupture process (e.g. Honda *et al.*, 2004, Sekiguchi *et al.*, 2006). In our routine analysis, different bandpass filters are used according to the magnitude M_{MA} , estimated by the JMA (2005): 0.02-0.05, 0.01-0.05 and 0.005-0.02 Hz pass band are used for the magnitude ranges of $3.5 \le M_{MA} < 5.0$, $5.0 \le M_{MA} < 7.5$, and $7.5 \le M_{MA}$, respectively.

Green's function is computed by using the discrete wavenumber method developed by Saikia (1994), whose code is called FKPPROG. The velocity structure used for Green's function is shown in Table 1. This structure is the

same as the NIED F-net routine, and is based on Ukawa et al. (1984) for the shallower parts, Fukao (1977) for the middle deep parts, and iasp91 (Kennett and Engdahl, 1991) for the deeper parts. Variance reduction (VR), which indicates the fit between observed and synthetics, is computed by the following equation:

$$VR(\%) = 100 \times \sum_{i} w_{i} \int \left(1 - \frac{(s_{i}(t) - o_{i}(t))^{2}}{|o_{i}(t)|^{2}}\right) dt$$

where $s_i(t)$ and $o_i(t)$ are synthetic and observed waveforms at station i, respectively. w_i is a weighting function proportional to the hypocentral distance (Fukuyama and Dreger, 2000). We select moment tensors whose VR is greater than 50%. The fixed epicenter locations are taken from the unified JMA hypocenter catalog. The horizontal distance from the epicenter to the station is rounded to the nearest interval of 5 km. The variance reduction approach has also been applied to focal depths from 5 km with an interval of 3 km. The results with maximum variance reduction are taken as the optimal focal depth.

In order to analyze moment tensor solutions for small earthquakes as far as possible, our new approach differs from that of the NIED F-net routine in the following respects: (a) The horizontal distance to the station is rounded to the nearest interval of 1 km. (b) The variance reduction approach was also applied to focal depths from 2 km with an interval of 1 km (Matsumoto *et al.*, 2006). (c) Filters with passbands of 0.02-0.05 Hz are used for a magnitude range of $3.0 \le M_{MA} < 5.0$. We can implement this advanced approach in the NIED F-net routine with a slight increase in processing time.

3. Results

From March 20 until June 30, 99 events with a magnitude exceeding 3.5 were listed in the JMA Catalog. In the NIED F-net routine, whose *VR* exceeded 50 percent and more than 3 stations were used, we obtained the moment tensors of 62 events. On the other hand, in this study, 272 events with a magnitude exceeding 3.0 were listed in the JMA Catalog,

	I	Main-shock		Largest aftershock			
	F-net	USGS	Harvard	F-net	USGS	Harvard	
	(this study)			(this study)			
Date/time (UT)	200	05/03/20 01:53	3	200	05/04/19 21:11		
Depth (km)	10	13	12	8	18	17	
Mw	6.6	6.5	6.6	5.6	5.5	5.4	
Strike (degree)	213/122	124/34	122/32	132/226	57/149	42/132	
Dip (degree)	83/83	87/89	89/82	79/69	78/81	88/86	
Rake (degree)	-173/-7	1/177	8/179	-21/-168	171/12	176/2	

Table 2. Source parameters obtained for the mainshock and the largest aftershock.

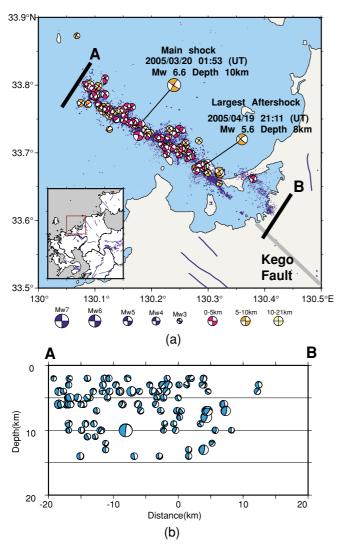


Fig. 2. Focal mechanism distribution. (a) Map-view. The color of the moment tensors indicates the depth, while blue dots indicate epicenters and bule lines indicate active faults. (b) The cross section along A-B in (a).

and the moment tensors of 101 events were determined. In the moment magnitude range of $Mw \le 4.0$, VR in this study went up 5 percent from that of F-net routine processing. In the moment magnitude range of Mw > 4.0, VR in this study was almost equal to that of F-net routine processing.

The spatial distribution of the moment tensors that were determined in this study is shown in Fig. 2. We successfully

obtain the moment tensors of the main-shock with very good variance reduction. The centroid depth error of these moment tensors is estimated to within 5 km (Fukuyama and Dreger, 2000), while the epicenter error of these events listed in the JMA Catalog is estimated to within 0.5 km. We use the fixed epicenter in the JMA Catalogue and the horizontal distance to the station rounded to the nearest interval of 1 km, then the location error of the horizontal direction is estimated to within 1 km. One of the nodal planes of the moment tensor solution is consistent with aftershock distribution. The source parameters for the main-shock and the largest aftershock are shown in Table 2 and are similar to those listed in the USGS CMT and Harvard CMT Catalogs. The focal mechanisms of the main-shock and the largest aftershock are of the strike-slip type.

We determined many moment tensor solutions of intermediate-sized aftershocks around the focal region. The Harvard CMT catalog effectively covers earthquakes with Mw > 5.3, the NIED F-net routine MT catalog those with Mw > 3.8 (Kubo *et al.*, 2002), and this study those with Mw > 3.5.

However, we were unable to determine the moment tensors of 35 events which occurred within 4 hours of the mainshock in the NIED F-net routine, and of the 90 events in this study. When a large earthquake occurs, it is difficult to perform waveform inversion with the swarm of the aftershock. In the case of the 2004 Mid Niigata earthquake (M_{JMA} 6.8) and the 2003 Tokachi-Oki earthquake (M_{JMA} 8.0), we were also unable to determine the moment tensors of aftershocks for 3 or 4 hours after the main-shock.

As a result, we obtained 101 moment tensors solutions distributed around the focal region.

4. Discussion

We thoroughly determined the moment tensor solutions of the main shock and intermediate-sized aftershocks around the focal region. As shown in Fig. 2, the focal mechanism of aftershocks obtained in this study is mainly of the strike-slip type. The trend of the epicenter distribution may correspond to the strike direction of the focal mechanism. The alignment of the epicenters in the rupture zone of the main-shock is oriented between N110°E and N130°E, which is close to the strike of the main-shock's moment tensor solutions (NIED: N122°E; USGS: N124°E; Harvard: N122°E). In the southeastern part of the focal region, we determined only two moment teonsor solutions. In this region, Imanishi

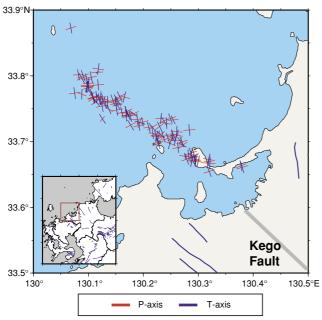


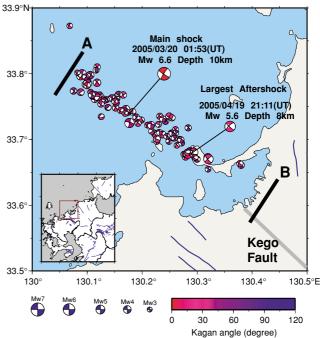
Fig. 3. Distributions of horizontally projected P-axis (pseudo P-axis) and T-axis (pseudo T-axis) azimuth directions. Red is P-axis, Blue is T-axis.

et al. (2006) determined the double-difference earthquake locations and focal mechanism solutions, these focal mechanisms are comparable with our study.

Figure 3 represents the P- and T-axes distribution of all moment tensor solutions. These P-axes with the direction from E-W to ENE-WSW are comparable with those of the stress field (Seno, 1999) in this region. However, there are some aftershocks with P- and T-axes being different from those of the main shock.

To show the similarity between the focal mechanisms of the main-shock and the aftershock, we used Kagan's angle (Kagan, 1991) defined as the minimum angle of rotation from one mechanism to the other in 3-D, which varies from 0 to 120°. Figure 4 represents the Kagan's angle distribution and the distributions of coseismic slip estimated by using strong ground motion data (Sekiguchi *et al.* 2006); the focal mechanisms colored red or pink are similar to the main-shock with a Kagan's angle of less than 30 degrees. In the inflection-point and the eastern part of the epicenter distribution, there are aftershocks with a Kagan's angle above 45 degrees.

The large aftershocks with Mw > 5.0 occurred in the area within 1 m of the main shock's slip. The intermediate-size aftershocks with a Kagan's angle less than 15 degrees occurred in the area within 1 m of the main shock's slip except the aftershocks in the largest slip area with 2 km depth. Nishimura *et al.* (2006) also estimated coseismic slip distribution with ground deformation date observed by GPS and InSAR. They proposed that large slip area is concentrated at depths of 0 to 10 km in an along-strike range from the hypocenter to 10 km east-southeast of it. The intermediate-size aftershocks with a Kagan's angle less than 15 degrees also occurred in the area within 1 m of the main shock's slip except the aftershocks in the largest slip area with 4 km depth.



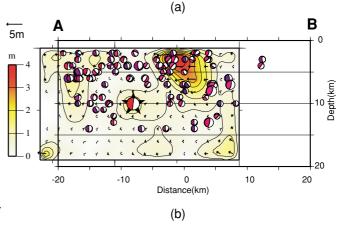


Fig. 4. Focal mechanism distribution. Color of moment tensors indicates the Kagan's angle. (a) Map-view. Blue lines indicate an active fault. (b) The cross section along A-B in (a) with distributions of coseismic slip obtained for the 2005 West Off Fukuoka Prefecture Earthquake by Sekiguchi et al. (2006). Arrows show the amplitude and directions of slip. Contour interval is 0.5 m.

On the other hand, the clear inflection-point of epicenter distribution is near 130.28°E. In the eastern part of the focal region, the alignment of the epicenters is oriented between N125°E and N145°E, which is close to the strike of the largest aftershock's moment tensor solutions. It occurred within the northwestern extent of the Kego fault, and the strike direction in the focal mechanism is consistent with that of the Kego fault.

It is possible that the largest aftershock occurred in another fault plane. For example, Ito *et al.* (2006) proposed that three different faults exist in this focal region by analyzing CMT solutions with velocity and acceleration seismograms of the F-net broadband seismometers and Hi-net tiltmeters. It seems that these CMT solutions indicate a more accurate mechanism and centroid location due to the use of additional waveform data for the analysis, meaning that

fault model is also supported by the strike direction of the focal mechanism of the main-shock and large aftershocks determined in this study. On the other hand, assuming the extended approach in the NIED F-net routine can determine more moment tensor solutions with magnitudes exceeding 3.0 all over the focal region, then it is possible that the stress field inferred from the distribution of the *P*- and *T*-axes and the spatial distribution of the moment tensor solutions with intermediate-sized aftershocks determined in this study are comparable with that fault model.

The NIED F-net routine processing is basically automated today. By implementing this extended approach in the NIED F-net routine, we will automatically determine the spatial distribution of many moment tensor solutions like this study in semi-real-time to understand the regional tectonic stress fields and source mechanisms of large earthquakes and so on.

5. Conclusion

We analyzed the moment tensor solutions with a magnitude exceeding 3.0 listed in the JMA Catalog, with the extended method of the NIED F-net routine processing and consequently obtained 101 moment tensor solutions. The main-shock's moment tensor solution in our study is comparable with the USGS CMT catalog and Harvard CMT catalog. We determined many moment tensor solutions of intermediate-sized aftershocks around the focal region. These moment tensor solutions represent basic and important information concerning earthquakes in investigating regional tectonic stress fields, source mechanisms, fault model and so on.

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