Indian plate kinematic studies by GPS-geodesy

E. C. Malaimani¹, James Campbell², Barbara Görres², Holger Kotthoff², and Stefan Smaritschnik²

¹National Geophysical Research Institute (NGRI), Hyderabad 500007, India ²Geodätisches Institut der Universität Bonn (GIUB), Nußallee 17, D-53115 Bonn, Germany

(Received December 28, 1999; Revised May 19, 2000; Accepted May 22, 2000)

Amongst various space techniques, GPS-geodesy has proved to be an indispensible tool worldwide for geodynamics and the determination of parameters governing seismotectonics. To study the Indian plate kinematics, an IGS station has been established in September 1995 at NGRI, Hyderabad, India, under the Indo-German collaborative research programme. During the first off-line data transmisssion phase, about one to two days per week of observations were taken. The data analysis has been carried out at the University of Bonn using the Bernese Software version 4.0. The Hyderabad data have been processed together with the data from 8 other selected IGS stations in and around the Indian plate. In our preliminary analysis, which concentrated on the ITRF96 frame-consistent part of the data set, we found the motion of the Indian plate with respect to the Eurasian plate to be significantly smaller than the model rate, i.e. 3.7 cm/y for Hyderabad instead of 4.5 cm/y in the NUVEL-1A plate model. We detected no significant motion between Hyderabad and Bangalore, lying 500 km apart in N-S-direction on the Indian plate.

1. Introduction

The lofty Himalayas, which stretch 2900 km along the border between India and Tibet are the visible creations of plate tectonic forces. The most dramatic effect of the moving Indian plate on its continental region is manifested at the Himalayan collision boundaries. The high elevations of the Himalaya and Tibet result from the continuing collision between India and Asia, which started more than 60 million years ago (Molnar, 1984). Global plate models (DeMets *et al.*, 1994) predict that approximately 5.0 cm/y of northward directed convergence is taken up between India and Eurasia.

Models of current plate motions have been estimated for almost 20 years using spreading rates at mid ocean ridges, transform fault azimuths and plate boundary earthquake slip vectors (DeMets *et al.*, 1990). Early in the development of plate tectonics it was recognized that plate boundaries in continental areas are substantially wider than those in oceanic plates.

From geological and seismic studies of the slip rate of faults in Asia, it is believed that approximately one third of the present day convergence rate between India and Asia is responsible for the shortening and uplift as well as the moderate seismicity of the Himalayas. To accommodate this convergence, studies by Avouac and Tapponnier (1993), England and Molnar (1997) and Molnar (1987) have explained that some part of the India and Eurasia collision is taken up by strike slip faulting in Asia, with corresponding eastward displacement or extrusion of Tibet and Southern China. Crustal thickening evidenced in the ~5 km average elevation of the Tibetan plateau and a crustal thickness of ~70 km represents the convergence clearly (Larson *et al.*, 1999). Geody-

namic models of this region heavily depend on these factors whether it is crustal thickening or lateral extrusion (England and Houseman, 1986; Royden *et al.*, 1997).

Many of these questions could be resolved if precise measurements of surface deformation throughout the plate boundary region were available. Amongst various space techniques, GPS-Geodesy has proved to be an indispensible tool worldwide for geodynamics and the determination of parameters governing seismotectonics as well as to quantify continental deformation to a precision of a few mm/y. With the successful operation of the International GPS Service (IGS) and its global network and as a participant in the GPS global campaign we will focus this paper on the modalities used to study the Indian plate kinematics using GPS.

2. Observations and Data Processing

In September 1995, a GPS tracking/reference station designed to become an operational IGS permanent station has been established at NGRI, Hyderabad, by installing a TurboRogue SNR-8000 receiver under the Indo-German collaborative research programme. Since then, an off-line data transfer to the University of Bonn has been implemented and data from 1995 onwards is being archived and available at NGRI, Hyderabad as well. Recently, an ftp-link has been established for daily data transfer, which will become operational early in 2000.

The GPS processing using the Bernese Software version 4.0 (Rothacher and Mervart, 1996) was carried out at the University of Bonn. A first data set for the data up to January 1997 was prepared and a first network solution was carried out to estimate baselines between Hyderabad and other selected stations which are located in and around the Indian plate.

Our aim in this study was to include all the data so far acquired at Hyderabad from September 1995 to August 1998,

Copy right© The Society of Geomagnetism and Earth, Planetary and Space Sciences (SGEPSS); The Seismological Society of Japan; The Volcanological Society of Japan; The Geodetic Society of Japan; The Japanese Society for Planetary Sciences.

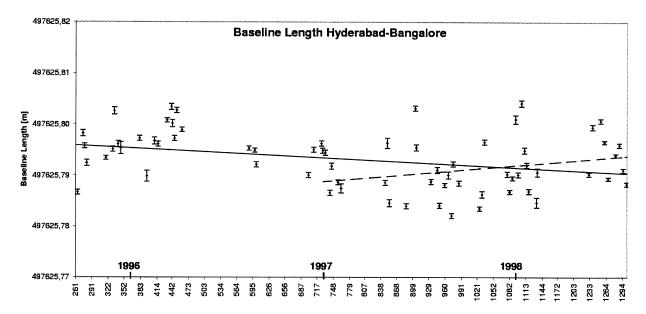


Fig. 1. Temporal baseline length evolution for Hyderabad-Bangalore.

Table 1. Stations used in our solutions.

Site	Latitude	Longitude
Bangalore	13.0219	77.5628
Irkutsk	52.2236	104.3090
Kitab	39.1355	66.8750
Lhasa	29.6594	91.0970
Shanghai	31.1030	121.1980
Wettzell	49.1393	12.8676
Hyderabad	17.1501	78.1982

but for the sake of computing time only one day of data per week was eventually included. We used the Bernese Processing Engine (BPE) to process the still large amount of data for 3 years from all the selected IGS stations (Bangalore, Yaragadee, Hartebeesthoek, Wettzell, Lhasa, Kitab, Irkutsk and Shanghai) (Table 1). No data decimation on the 30 sec sampling rate was made, i.e. all of the original data were retained for the processing.

For the network solution, we developed the following strategies (Campbell and Malaimani, 1998):

- i) IGS precise orbits: the accuracy is about ± 0.06 m
- ii) 4 tropospheric parameters/day
- iii) elevation cutoff at 20 degrees
- iv) no ambiguity fixing
- v) reference frames: for our study we have adopted the ITRF94 reference frame (Boucher *et al.*, 1996) to analyse the data from 1995 to 1996. By January 1997, the ITRF96 reference frame was available to the global users by IERS, hence we used ITRF96 site coordinates and velocities to process the data from 1997 to 1998 (Boucher *et al.*, 1998). During processing, we realized that the small shifts and rotations between successive ITRF realizations could not be neglected.

Therefore we carried out two solutions, one covering the entire period of September 1995 to August 1998 and a second one concentrating only on the data of January 1997 to August 1998, when the ITRF96 was valid for the IGS final orbits.

vi) fiducial free strategy: in order to minimize the constraints and gain an independend insight in the integrity of our results we decided to fix only the station of Wettzell on the Eurasian plate. In this way, our network solution provides the motions of all of the other stations referred to the stable part of Eurasia and the IGS orbital frame.

3. Results and Discussion

Using the above strategies, the network solution was carried out. We estimated the baseline vectors and the rate of change of baseline lengths from Hyderabad to the other IGS stations, as well as velocity vectors for all stations involved in the analysis (relative to Wettzell).

The overall result of this analysis is that the estimated baseline lengths between Hyderabad and other IGS stations beyond the Himalayas are indeed shortening. They clearly indicate the effect of the northward trend of he motion of the Indian plate.

The effect of the transition between the ITRF94 and ITRF96 frames can be seen quite clearly in the baseline length evolution plots (Figs. 1 to 3) and even more so in the coordinate evolution plots (not shown here). Before having established proper transformations to account for this transition we decided to confine our vector solution to the consistent data period from January 1997 to August 1998 of the ITRF96. The estimated baseline rates are shown for both the entire period as well as for the ITRF96 consistent period (Table 2).

When applying the linear fit to the baseline length results of the individual epochs, several common features emerged throughout the estimation:

1) The formal errors of the individual epoch solutions are much smaller than the rms errors from the linear fit, which

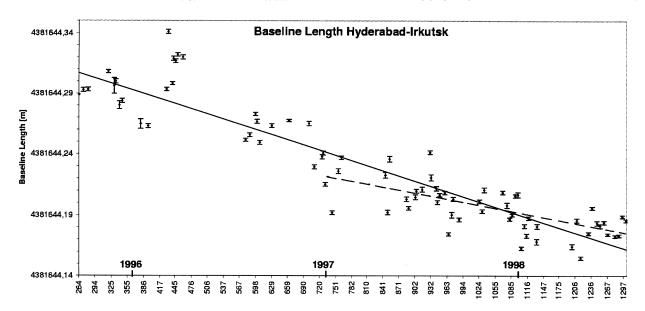


Fig. 2. Temporal baseline length evolution for Hyderabad-Irkutsk.

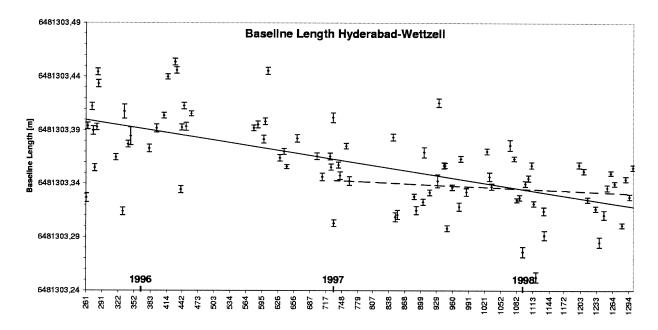


Fig. 3. Temporal baseline length evolution for Hyderabad–Wettzell.

attests the presence of unmodelled errors, probably of tropospheric origin.

- 2) The increase of both types of errors follows the empirical error relation $\sigma^2 = A^2 + B^2 \cdot L^2$ with A = 0.5 cm and $B = 10^{-8}$ (Lichten, 1990). In fact, if we discard the isolated long baseline to Yaragadee the coefficient B comes very close to $0.6 \cdot 10^{-8}$.
- 3) The difference in the rates derived from the two solutions highlights the importance of the use of a consistent frame in the analysis of large networks.

Looking at the individual results from the 1997 to 1998 data, there are several significant insights to be gained:

HYDERABAD to BANGALORE (Fig. 1): The temporal change of this Indian baseline of 500 km length is only +0.3

cm/y with an rms scatter of ± 0.5 cm, which indicates a high level of inner plate stability. This value provides a first constraint on the stability of the Indian plate in its central and southern part.

HYDERABAD to IRKUTSK (Fig. 2): This baseline stretches far north into the more stable part of the Eurasian plate and shows a consistent shortening rate of -3.1 cm/y. With an rms about the linear fit of ± 1.5 cm and no systematic variations over time, this determination holds a high degree of reliability.

HYDERABAD to LHASA: The low rate of shortening of -0.6 cm/y confirms the heavy involvement of this part of southern Tibet in the Indo-Eurasian collision zone.

HYDERABAD to WETTZELL (Fig. 3): This very long

Table 2. Estimated baselines from Hyderabad and their rates of change for the total time span (September 95–August 98) and the ITRF96 period (January 97–August 98).

From Hyderabad to	baseline length at ref. epoch 6th March 1997	ITRF94/96 rate of change [cm/y]	ITRF94/96 rms scatter [cm]	ITRF96 rate of change [cm/y]	ITRF96 rms scatter [cm]
Bangalore	497,625.797	-0.2	±0.5	+0.3	±0.5
Shanghai	4,459,795.268	-1.2	± 2.3	-2.2	± 2.0
Lhasa	1,856,738.483	-0.9	± 0.9	-0.6	± 0.9
Irkutsk	4,381,644.341	-5.0	±1.9	-3.1	±1.5
Kitab	2,640,353.581	-3.0	±1.1	-1.9	±1.2
Wettzell	6,481,303.421	-2.9	±3.2	-0.3	± 3.2
Yaragadee	6,208,968.830	+2.1	±9.3	+4.0	±9.9

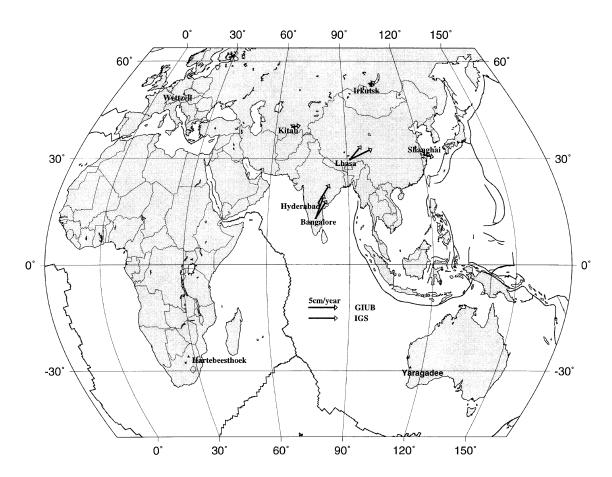


Fig. 4. Velocity vectors from GIUB (Bonn) and IGS solution.

baseline of 6500 km shows a very good repeatability with an rms of only $\pm 3 \text{ cm}$. It does not show any significant change in the period of the ITRF96.

In our post-fit analysis we obtained a preliminary velocity solution in the NUVEL-1 frame, i.e. in a No-Net-Rotation (NNR) reference frame (DeMets *et al.*, 1994) by applying linear fits to the site-coordinate time series. The NUVEL-1 frame is also used to constrain the IGS and ITRF solutions (Boucher *et al.*, 1994, 1996). Because our intention was to

study the Indian plate motion with respect to a Eurasia-fixed frame, we transformed the velocities to this frame by subtracting the Eurasian (NUVEL) plate rotation. To compare our solution with the IGS solution for the ITRF96 we show both vectors for all points except Hyderabad (Fig. 4). For the Eurasian stations of Kitab and Irkutsk as well as for Bangalore the agreement with IGS is quite good whereas larger differences appear at the stations of Lhasa and Shanghai.

From our preliminary analysis, the relative motion of India

with respect to Eurasia turns out to be 3.7 ± 0.2 cm/y in NNE direction, which is significantly smaller than the NUVEL motion of 4.5 cm/y. This motion is virtually the same for both stations Hyderabad and Bangalore (Fig. 4).

4. Conclusions

We have analyzed 3 years of GPS data from September 1995 to August 1998 involving the new site of Hyderabad on the Indian plate. Although our frame-consistent solution only covers 1.5 years of data we were able to determine the rate of movement of the Indian plate. The fiducial free solution also provides an unconstrained insight in the kinematics of the other sites included in the analysis. The linear fits of the baseline lengths yield highly significant shortenings between India and Eurasia of up to -3.1 cm/y. The Indian baseline Hyderabad–Bangalore shows no significant deformation over this relatively short time span.

The site velocities from our preliminary solution (with only Wettzell fixed) in general show a good agreement with the NUVEL-1A geological averages, although our GPS based rates for the two sites on the Indian plate of 3.7 cm/y (relative to Eurasia) are significantly smaller than the corresponding NUVEL rates of 4.5 cm/y. However, the reduced rate of 3.7 cm/y appears to be in perfect agreement with the recently revised velocity of 3.7 cm/y in the plate model described by Gordon *et al.* (1999).

Acknowledgments. The authors are thankful to Dr. Harsh K. Gupta, Director, NGRI, HYDERABAD for his continued support for this project right from the start. This project is being funded by the Indo-German Collaborative project under International Scientific CSIR-DLR bilateral exchange programme.

References

Avouac, J. P. and P. Tapponnier, Kinematic model of active deformation in Asia, Geophys. Res. Lett., 20, 895–898, 1993.

- Boucher, C., Z. Altamimi, M. Feissel, and P. Sillard, Results and analysis of the ITRF94. *IERS Technical Note 20*. Paris, 1996.
- Boucher, C., Z. Altamimi, and P. Sillard, Results and analysis of the ITRF96, *IERS Technical Note 24*, Paris, 1998.
- Campbell, J. and E. C. Malaimani, Estimation of very long baselines by GPS geodesy for Indian plate kinematic studies, Technical Report No. NGRI-98-SEIS-235, 1998.
- DeMets, C., R. G. Gordon, D. F. Argus, and S. Stein, Current plate motions, Geophys. J. Int., 101, 425–478, 1990.
- DeMets, C., R. G. Gordon, D. F. Argus, and S. Stein, Effect of recent revisions to the geomagnatic reversal time scale on estimates of current plate motions, *Geophys. Res. Lett.*, 21, 2191–2194, 1994.
- England, P. C. and G. A. Houseman, Finite strain calculation of continental deformation, 2, Comparison with the India-Asia collision zone, *J. Geophys. Res.*, 91, 3664–3676, 1986.
- England, P. C. and P. Molnar, Active deformation of Asia: From kinematics to dynamics, *Science*, 278, 647–650, 1997.
- Gordon, R. G., D. F. Argus, and M. B. Heflin, Revised estimate of the angular velocity of India relative to Eurasia, Supplement to EOS Transactions, AGU, Vol. 80, No. 46, p. F273, 1999.
- Larson, K. M., R. Bürgmann, R. Bilham, and J. T. Freymueller, Kinematics of the India—Eurasia collision zone from GPS measurements, *J. Geo*phys. Res., 104, 1077–1093, 1999.
- Lichten, S. M., High accuracy global positioning system orbit determination: Progress and prospects, in *Global Positioning System: An Overview, IAG Symposium No. 102*, edited by Y. Bock and N. Leppard, pp. 146/164, Springer, New York, 1990.
- Molnar, P., Structure and tectonics of the Himalaya: Constraints and implications of geophysical data, *Annu. Rev. Earth Planet. Sci.*, **12**, 489–518, 1984.
- Molnar, P., Inversion of profiles of uplift rates for the geometry of dipslip faults at depth, with examples from the Alps and Himalaya, *Ann. Geophys.*, **5**, 663–670, 1987.
- Rothacher, M. and L. Mervart (Eds.), Bernese GPS Software Version 4.0 Reference Manual, Berne, 1996.
- Royden, L. H., B. C. Burchfiel, R. W. King, E. Wang, Z. Chen, F. Shen, and Y. Liu, Surface deformation and lower crustal flow in eastern Tibet, *Science*, 276, 788–790, 1997.

E. C. Malaimani, J. Campbell (e-mail: jcampbell@picasso.geod.unibonn.de), B. Görres, H. Kotthoff, and S. Smaritschnik