

Geochronological constraint on the brittle-plastic deformation along the Hatagawa Fault Zone, NE Japan

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K-Ar ages and fission-track ages of granitic rocks in the Hatagawa Fault Zone (HFZ), NE Japan were measured to examine the cooling history of the HFZ. The HFZ is an NNW-SSE trending fault zone in the Cretaceous granitic rocks, and consists of a conspicuous cataclasite and two types of mylonite with a sinistral sense of shear. The cataclasite zone is NNW-SSE trending and continuous over at least 40 km with a maximum thickness of 100 m. One type of mylonite is low-T mylonite, which is mainly developed for a length of 6 km along the HFZ. The other is high-T mylonite, which is widely distributed in the HFZ. Most of K-Ar ages of hornblende and biotite from granitic rocks are about 110 Ma and show no obvious differences along the strike of the HFZ or among different granite bodies. This implies that the granitic rocks in the HFZ have a similar cooling history and cooled rapidly from closure temperature of hornblende to that of biotite. Zircon fission-track analysis shows little possibility of reheating of the granitic rocks. This supports the formation of cataclasite and mylonite during the cooling of the granitic bodies. Fission-track ages of zircon and apatite from the samples in and near the areas where the low-T mylonite is developed are older than those for other areas. Infiltration of near-surface derived water into the low-T mylonite after plastic deformation may account for the accelerate cooling of granitic bodies.

Key words: Hatagawa Fault Zone, brittle-plastic deformation, K-Ar ages, fission-track ages.

1. Introduction

The thermal history along a fault strongly affects faulting history, and fault displacement often governs the thermal history along the fault (e.g. Stöckhert *et al.*, 1999). Recent studies of the Hatagawa Fault Zone (HFZ), NE Japan, have revealed that (1) cataclasite and mylonite formed within a limited period between 126 ± 6 and 98.1 ± 2.5 Ma (Tomita *et al.*, 2002), (2) repeated episodes of plastic deformation and fracturing occurred in the mylonite zone (Takagi *et al.*, 2000; Tomita *et al.*, 2002; Shigematsu *et al.*, 2004), (3) cataclasite formed at temperature above 220°C (Tomita *et al.*, 2002) and (4) the lowest estimation of deformation temperature of mylonite is 300°C (Shigematsu *et al.*, 2003). As no obvious discontinuities of time and temperature are found between cataclasite and mylonite formation, the fault rocks along the HFZ passed through the brittle-plastic transition zone (Takagi *et al.*, 2000; Tomita *et al.*, 2002; Shigematsu *et al.*, 2004). Shigematsu *et al.* (2003) deduced that heterogeneous distribution of plastic deformation in the brittle-plastic transition along the HFZ resulted in a significant

stress concentration and the nucleation of large earthquakes. The HFZ enables us to examine the relation between the cooling history and the heterogeneity of plastic deformation in the brittle-plastic transition, especially with the nucleation zone of the large earthquakes. K-Ar dating and fission-track dating of granitic rocks were, therefore, performed in this study.

2. Hatagawa Fault Zone

The HFZ is at the eastern margin of the Abukuma Mountains, NE Japan (Fig. 1), and trends in the NNW-SSE direction (Watanabe *et al.*, 1953; Sendo, 1958). The South Kitakami belt is located to the east of the HFZ, and the Abukuma belt is to the west (Kubo and Yamamoto, 1990). The HFZ is located at the boundary of these two belts. The South Kitakami and the Abukuma belts around the HFZ mainly consist of magnetite-bearing and magnetite-free Cretaceous granitic rocks respectively. A conspicuous cataclasite zone with a maximum thickness of about 100 m extends continuously in the HFZ for at least 40 km (Fig. 2) (Tomita *et al.*, 2002). Mylonite zones with a sinistral sense of shear are heterogeneously distributed with a maximum thickness of 1 km along the HFZ (Shigematsu and Yamagishi, 2002). The foliation strikes N-S to NE-SW and dips subvertically; the stretching lineation is subhorizontal. Quartz microstructures in the mylonite have been subdivided into two groups, microstructures A and B (Fig. 3) (Shigematsu and Yamagishi, 2002). Microstructure A is

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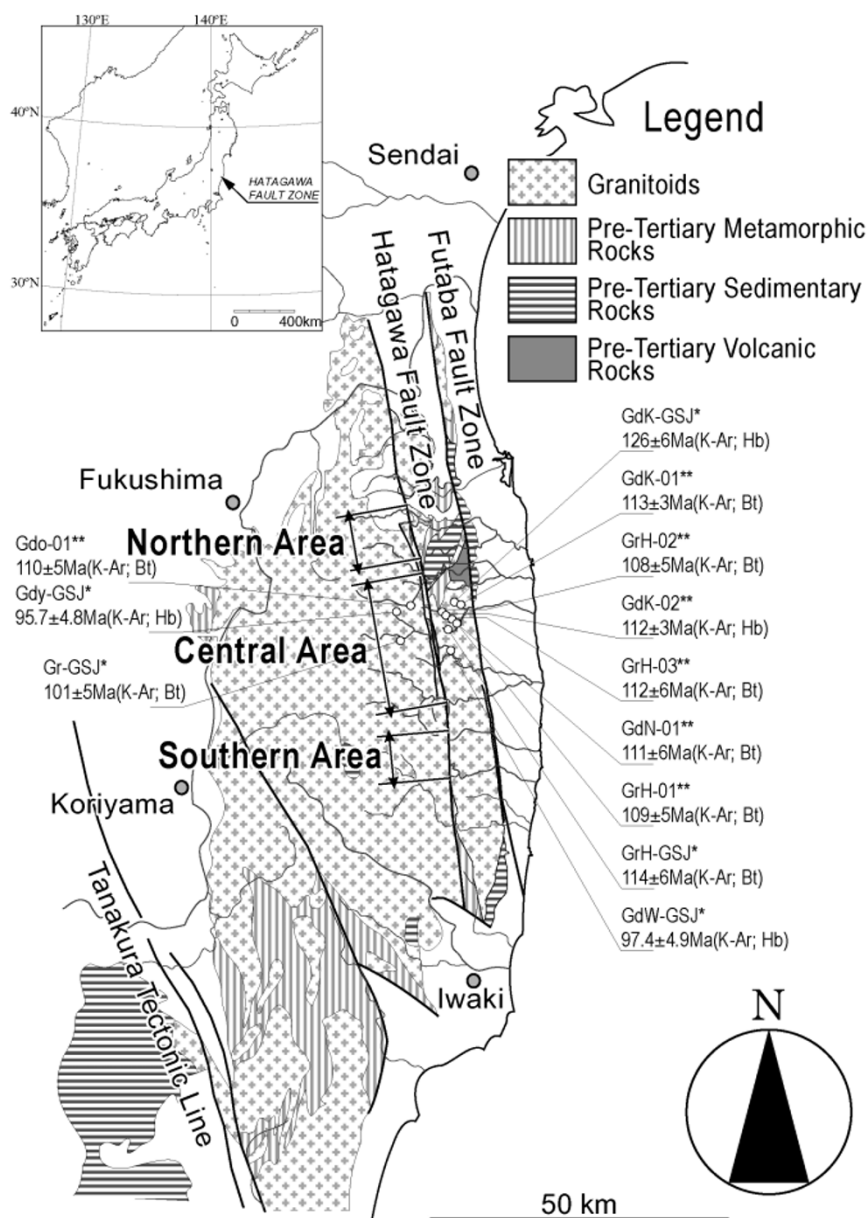


Fig. 1. Regional geological map of the Abukuma Mountains, NE Japan and K-Ar ages of samples away from the Hatagawa Fault Zone (Kubo *et al.*, 1990; Tomita *et al.*, 2002). * and ** represent the data from Kubo *et al.* (1990) and Tomita *et al.* (2002), respectively. Hb and Bt indicate hornblende and biotite, respectively. Three areas along the Hatagawa Fault Zone, northern area, central area and southern area are shown in Fig. 2.

characterized by quartz grains exhibiting core and mantle structure (White, 1976) and a type I crossed girdle fabric pattern resulting from subgrain rotation dynamic recrystallization. Microstructure B is characterized by quartz grains with a Y-maximum fabric pattern caused by both progressive subgrain rotation and grain boundary migration. The mylonite with quartz microstructure A is mainly developed for a length of 6 km along the strike of the HFZ, while the mylonite with quartz microstructure B is widely distributed in the HFZ (Shigematsu *et al.*, 2003). Shigematsu and Yamagishi (2002) calculated deformation temperatures of these mylonites by the thermometer after Stormer (1975), and Shigematsu *et al.* (2003) recalculated them by the two-feldspar thermometer (Whitney and Stormer, 1977). The recalculated temperatures generally range from 300°C to 360°C for mylonite with quartz microstructure A and from

340°C to 480°C for the mylonites with quartz microstructure B. Some mylonites with quartz microstructure A exhibit deformation temperatures ranging between 340°C to 480°C. In this paper, the mylonites which exhibit deformation temperature ranging from 300°C to 360°C are called low-T mylonites, and those ranging from 340°C to 480°C are high-T mylonites.

3. Geochronological Results

Nine samples were collected for K-Ar and fission-track datings to constrain the cooling history of the HFZ (Fig. 2). These samples were taken from undeformed granitic rocks near the mylonite zone or weakly deformed granitic rocks in the mylonite zone. K-Ar ages of 8 samples and fission-track ages of 4 samples were determined in this study (Tables 1 and 2). Some geochronological results in the region of

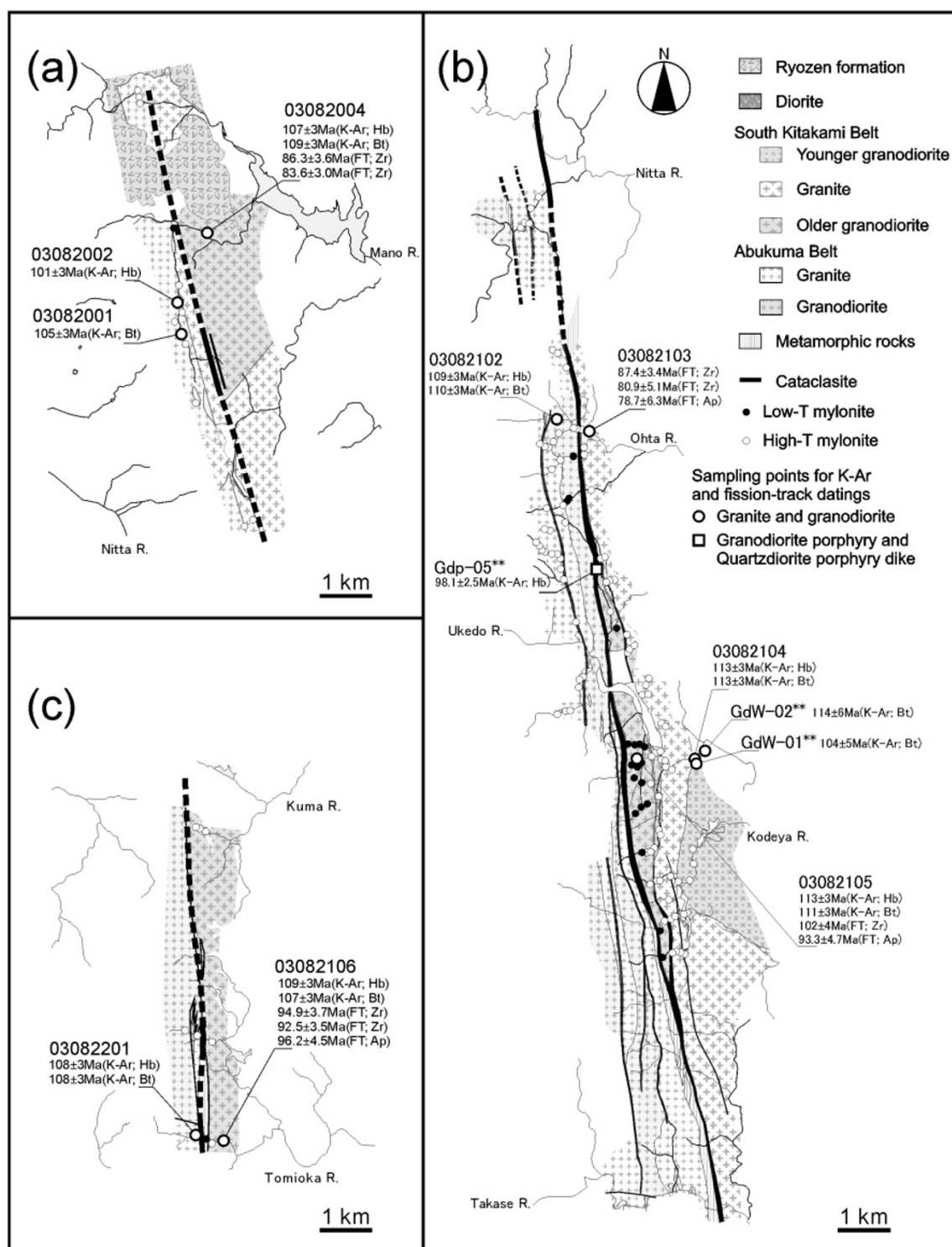


Fig. 2. Simplified geological map along the Hatagawa Fault Zone, showing distribution of low-T and high-T mylonites and K-Ar and fission-track ages. Compiled from Kubo *et al.* (1994), Yanagisawa *et al.* (1996) and Shigematsu and Yamagishi (2002). Ages with * and ** represent the data from Kubo *et al.* (1990) and Tomita *et al.* (2002), respectively. Hb, Bt, FT, Zr, Ap indicate hornblende, biotite, fission-track, zircon and apatite, respectively. (a) Northern area, (b) Central area and (c) Southern area.

the Abukuma Mountains have previously been reported by Kubo *et al.* (1990) and Tomita *et al.* (2002). Here, we combine these published results with those of our study.

K-Ar ages of hornblende from samples of granitic rocks in the HFZ range from 113 ± 3 to 101 ± 3 Ma, and those of biotite show a range from 114 ± 3 to 104 ± 5 Ma (Fig. 2). Most samples of granitic rocks newly measured in this study

show almost similar K-Ar ages of ca. 110 Ma, allowing for the error range of the dating, except for a few samples. Six samples have yielded both K-Ar ages of hornblende and biotite from a single sample (03082004, 03082102, 03082104, 03082105, 03082106 and 03082201). In these samples, the difference between K-Ar ages of hornblende and biotite is not significant. The samples from the western

Table 1. K-Ar ages from the granitic rocks in the Hatagawa Fault Zone (HFZ). * and ** represent the data from Kubo *et al.* (1990) and Tomita *et al.* (2002), respectively. Hb and Bt indicate hornblende and biotite, respectively. Samples from this study were analyzed at Geospace Science Co., Ltd.

Sample No.	Material Analyzed	⁴⁰ Ar rad (x10 ⁻⁵ scc/g)	⁴⁰ Ar rad (%)	K (%)	Isotopic Age (Ma)
Granitic rocks on eastern side of the HFZ					
* GdW-GSJ	Hb	0.430	86.8	1.14	97.4 ± 4.9
		0.461	87.2	1.15	
** GdW-01	Bt	1.64	93.3	4.02	104 ± 5
		1.72	93.5	4.02	
** GdW-02	Bt	1.63	96.1	3.50	114 ± 6
		1.58	95.4	3.52	
* GrH-GSJ	Bt	3.10	86.6	0.419	114 ± 6
		3.08	92.4	0.430	
** GdN-01	Bt	2.43	93.4	5.49	111±6
		2.46	93.6	5.46	
** GrH-01	Bt	1.92	93.8	4.33	109±5
		1.85	94.6	4.34	
** GrH-02	Bt	1.92	83.1	4.47	108 ± 5
		1.92	88.4	4.44	
** GrH-03	Bt	3.09	97.5	6.83	112±6
		3.06	97.4	6.81	
* GdK-GSJ	Hb	0.419	86.9	0.84	126±6
		0.430	92.4	0.84	
** GdK-01	Bt	2.84	93.2	6.22	113±3
		2.80	94.6	6.22	
** GdK-02	Hb	1.49	96.1	3.37	112 ± 3
		1.52	94.5	3.37	
03082004	Hb	0.385	85.5	0.91	107 ± 3
		0.392	83.2	0.91	
	Bt	3.12	95.0	7.12	109 ± 3
		3.11	96.4	7.12	
03082104	Hb	0.331	84.8	0.72	113 ± 3
		0.324	86.2	0.72	
	Bt	3.28	97.2	7.22	113 ± 3
		3.26	97.6	7.24	

Sample No.	Material Analyzed	⁴⁰ Ar rad (x10 ⁻⁵ scc/g)	⁴⁰ Ar rad (%)	K (%)	Isotopic Age (Ma)
03082105	Hb	0.273	81.8	0.61	113 ± 3
		0.278	82.2	0.61	
	Bt	3.00	96.4	6.72	111 ± 3
		2.98	96.1	6.70	
03082106	Hb	0.251	74.0	0.58	109 ± 3
		0.257	77.7	0.58	
	Bt	3.09	96.3	7.23	107 ± 3
		3.12	95.4	7.22	
Granitic rocks on western side of the HFZ					
03082001	Bt	2.86	96.6	6.85	105 ± 3
		2.87	96.2	6.85	
03082002	Hb	0.198	73.7	0.49	101 ± 3
		0.199	70.1	0.49	
03082102	Hb	0.294	87.9	0.67	109 ± 3
		0.293	82.4	0.67	
	Bt	2.80	95.3	6.33	110 ± 3
		2.78	96.1	6.35	
03082201	Hb	0.274	76.2	0.63	108 ± 3
		0.269	79.3	0.63	
	Bt	2.52	94.9	5.85	108 ± 3
		2.54	94.2	5.84	
** Gdo-01	Bt	2.52	97.9	5.73	110±5
		2.51	98.5	5.73	
* Gr-GSJ	Bt	2.52	90.9	2.55	101±5
		2.51	84.4	2.57	
* Gdy-GSJ	Hb	2.52	75.7	0.575	95.7±4.8
		2.51	88.9	0.586	
Dike intruding into the cataclastite zone					
** Gdp-05 (GdK)	Hb	0.097	58.2	0.25	98.1 ± 2.5
		0.099	56.8	0.25	

Table 2. Fission-track analysis of zircon and apatite from the granitic rocks in the Hatagawa Fault Zone. Two fission-track ages of zircon and one age of apatite were measured for each sample. Fission-track samples were analyzed at Kyoto Fission-track Co., Ltd. by following the procedure described by Danhara *et al.* (2003).

Sample		No. of crystals	ρ_s (cm^{-2})	N_s	ρ_i ($\times 10^6 \text{cm}^{-2}$)	N_i	ρ_d ($\times 10^4 \text{cm}^{-2}$)	N_d	r	$Pr(\chi^2)$ (%)	U (ppm)	Fission Track Age $\pm 1\sigma$ (Ma)	L μm	N	σ μm
03082004	Zr	30	9.18×10^6	4304	1.72	809	8.598	4402	0.512	28	190	86.3 ± 3.6	10.12	50	0.96
	Zr	23	1.06×10^7	5969	2.12	1196	8.874	4543	0.715	0	230	83.6 ± 3.0			
	Ap	30	6.05×10^5	662	1.11	1218	110.9	4435	0.678	24	12	$* 100 \pm 5$			
03082103	Zr	30	1.14×10^7	5296	2.14	993	8.680	4444	0.612	30	230	87.4 ± 3.4	10.42	50	0.93
	Zr	11	1.07×10^7	1585	2.22	329	8.892	4553	0.466	35	240	80.9 ± 5.1			
	Ap	29	2.32×10^5	240	5.46	564	111.1	4445	0.481	9	6	78.7 ± 6.3			
03082105	Zr	30	9.49×10^6	4849	1.51	773	8.625	4416	0.690	48	170	102 ± 4	10.47	50	0.78
	Zr	17	1.18×10^7	3751	2.51	794	8.880	4546	0.471	4	270	$* 79.2 \pm 3.4$			
	Ap	30	6.85×10^5	734	1.35	1452	111.0	4438	0.798	8	14	93.3 ± 4.7			
03082106	Zr	30	9.88×10^6	5493	1.70	945	8.652	4430	0.485	10	190	94.9 ± 3.7	10.62	50	0.67
	Zr	24	1.09×10^7	5607	1.97	1016	8.886	4549	0.713	94	210	92.5 ± 3.5			
	Ap	30	5.56×10^5	865	1.07	1659	111.0	4442	0.199	46	11	96.2 ± 4.5			

*. Ages excluded from discussion; Zr, zircon; Ap, apatite; ρ_s , spontaneous track density of a sample; N_s , number of tracks counted to determine ρ_s ; ρ_i , induced track density of a sample measured in DAP external detector; N_i , number of tracks counted to determine ρ_i ; ρ_d , induced track density of glass dosimeter NIST-SRM612 measured in a DAP external detector; N_d , number of tracks counted to determine ρ_d ; r , correlation coefficient between ρ_s and ρ_i ; $Pr(\chi^2)$, upper χ^2 tail probability corresponding to the observed χ^2 -statistics (Galbraith, 1981); L, mean length of confined fission tracks; N, number of measured tracks to determine L; σ , standard deviation for the length distribution of confined fission tracks; Thermal neutrons were irradiated in the pneumatic tube of JRR-4 reactor unit of Japan Atomic Energy Research Institute.

side of the cataclastic zone in the northern area show a hornblende K-Ar age of 101 \pm 3 Ma (03082002) and a biotite K-Ar age of 105 \pm 3 Ma (03082001), which are younger than those of most of the samples.

Some K-Ar ages previously reported away from the HFZ show different ages from 110 Ma (Fig. 1). In the eastern side of the HFZ (the South Kitakami belt), the sample GdW-GSJ gives a hornblende K-Ar age of 97.4 \pm 4.9 Ma. The sampling point of this sample is located in the center of the Kawabusa granodiorite. As two samples from the mar-

gin of this pluton exhibit biotite K-Ar ages of 104 \pm 5 Ma (GdW-01) and 114 \pm 6 Ma (GdW-02) (Fig. 2), these K-Ar ages would indicate the difference of cooling in this pluton (Tomita *et al.*, 2002). The sample GdK-GSJ away from the HFZ shows a hornblende K-Ar age of 126 \pm 6 Ma. Although this is the oldest K-Ar age in this region, K-Ar ages of two different samples (GdK-01, GdK-02) from the same pluton are about 110 Ma. In the western side of the HFZ (the Abukuma belt), Gdy-GSJ and Gr-GSJ far from the HFZ exhibit younger K-Ar ages of 95.7 \pm 4.8 Ma and 101 \pm 5 Ma

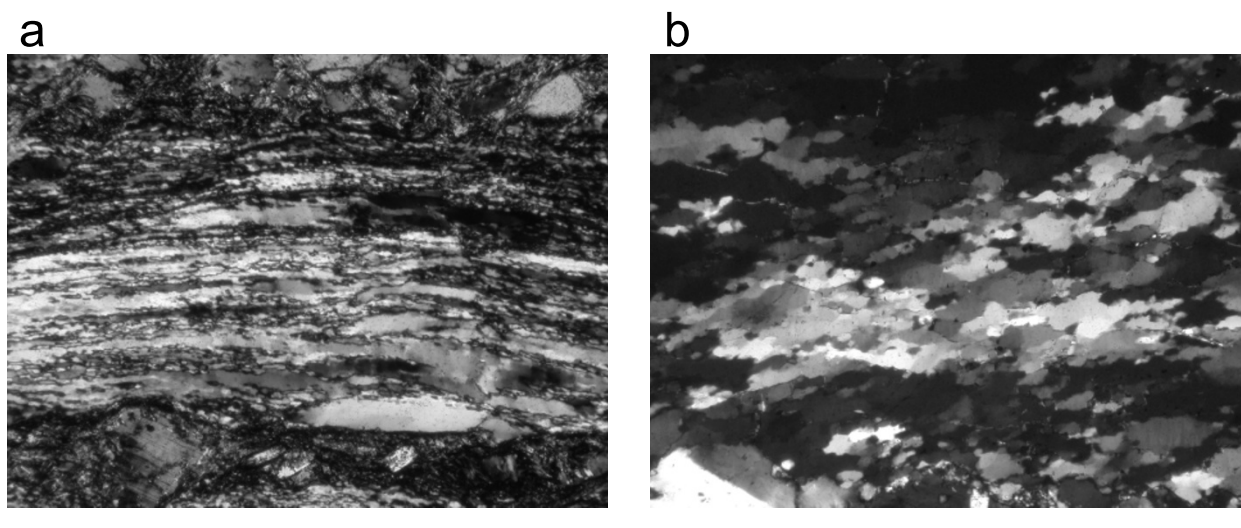


Fig. 3. Microstructure of the fault rocks from the Hatagawa Fault Zone. (a) Quartz microstructure of mylonite with microstructure A. Extremely elongated quartz grains with heavy undulatory extinction are surrounded by finer grains. (b) Quartz microstructure of mylonite with microstructure B. Aspect ratio of quartz grains does not depend on grain size. Scale bar is 0.1 mm.

respectively.

Fission-track ages of zircon and apatite were measured for four samples (Table 2). We excluded unreliable fission-track ages based on the χ^2 -test and distinction between spontaneous tracks and dislocations. Spontaneous tracks in apatite grains of sample 03082004 are hard to distinguish from dislocations. Fission-track ages of apatite obtained from this sample were older than those of zircon. Thus the apatite fission-track age obtained from this sample is not reliable. One result of zircon from sample 03082105 failed the χ^2 -test at the 5% significance level, and this fission-track age (79.2 ± 3.4 Ma) is quite different from the other zircon fission-track age (102 ± 4 Ma) of the same sample. These two fission-track ages are excluded from the following discussion (Table 2). Although one result of zircon from sample 03082004 failed the χ^2 -test at the 5% significance level, this fission-track age corresponds with the other fission-track age of zircon from the same sample within the error range. We consider this fission-track age to be meaningful. Excluding the unreliable ages, fission-track ages of zircon from 4 samples range from 102 ± 4 to 80.9 ± 5.1 Ma, and those of apatite show a range of 96.2 ± 4.5 to 78.7 ± 6.3 Ma. Average lengths of fission-track in zircon from 4 samples are 10.12 to 10.62 μm and do not show obvious shortening.

The fission-track ages of zircon and apatite from the samples in and near the area where the low-T mylonites are developed (03082105 and 03082106) are older than those from the others (03082004 and 03082103) (Fig. 2). Three samples of them have also yielded K-Ar ages of hornblende and biotite. The differences of fission-track age of zircon and K-Ar age of biotite from a single sample in and near the area where the low-T mylonites are developed are shorter than that of the high-T mylonites.

4. Discussion

The K-Ar systems of hornblende and biotite have different closure temperatures (Dodson and McClelland-Brown,

1985), and can be used to constrain the cooling history. Most of the granitic rocks in the HFZ show almost similar K-Ar ages within the error range of the dating except for a few samples, suggesting that granitic bodies cooled rapidly from closure temperature of hornblende to that of biotite. K-Ar dating along the HFZ reveals that the cooling history is almost the same along the strike of the HFZ, although granitic bodies represent multiple intrusions (Kubo and Yamamoto, 1990).

The closure temperatures of hornblende and biotite K-Ar ages are $510 \pm 30^\circ\text{C}$ (Dodson and McClelland-Brown, 1985) and $310 \pm 30^\circ\text{C}$ (Harrison *et al.*, 1985), and those of fission-track zircon and apatite are $280 \pm 30^\circ\text{C}$ and $110 \pm 10^\circ\text{C}$ (Stöckhert *et al.*, 1999), respectively. These temperatures can be compared to the estimated deformation temperatures of high-T mylonites (340°C to 480°C ; Shigematsu *et al.*, 2003), low-T mylonite (300°C to 360°C ; Shigematsu *et al.*, 2003), and cataclasite (220°C to 300°C ; Tomita *et al.*, 2002) (Fig. 4). These suggest that the formation of high-T mylonite started after approximately 110 Ma, and the deformation conditions were rapidly changed into those of low-T mylonite. The cataclasis of the HFZ had eventually terminated by 98.1 ± 2.5 Ma, as indicated by the K-Ar age of hornblende from an undeformed granodiorite porphyry dike intruded into the cataclasite zone (Tomita *et al.*, 2002). Average lengths of fission-track in zircon do not show obvious shortening. Tagami and Shimada (1996) proposed a zircon partial annealing zone between 230°C to 330°C . This suggests that the granitic rocks in the HFZ were not heated after the cooling to a temperature lower than 230°C and supports the formation of cataclasite and mylonite only during the cooling of the granitic bodies.

Shigematsu *et al.* (2003) estimated that plastic deformation was localized along the HFZ during the cooling of the granitic rock bodies from ca. 480 to 300°C , and the distributed area of low-T mylonite represents the final localized area of plastic deformation. Although the K-Ar ages of hornblende and biotite do not show a significant difference

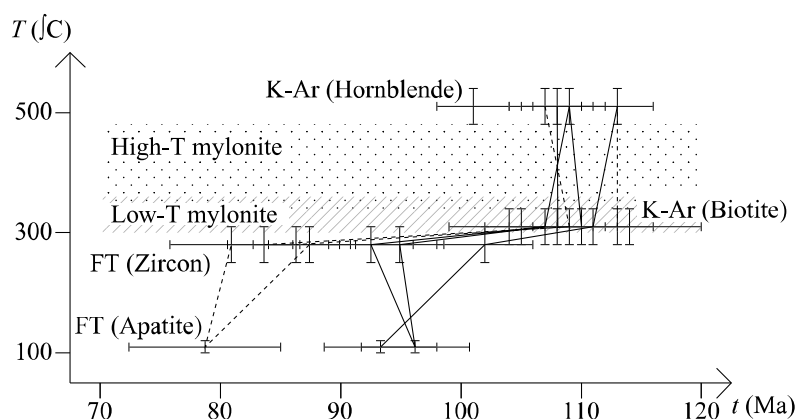


Fig. 4. Cooling history deduced from K-Ar and fission-track (FT) ages of granitic rocks in the Hatagawa Fault Zone. Connected lines between some ages mean that the ages were measured from the same sample. Solid and broken lines are cooling paths of low-T and high-T mylonites respectively. Closure temperatures of hornblende, biotite, zircon and apatite are $510 \pm 30^\circ\text{C}$ (Dodson and McClelland-Brown, 1985), $310 \pm 30^\circ\text{C}$ (Harrison *et al.*, 1985), $280 \pm 30^\circ\text{C}$ and $110 \pm 10^\circ\text{C}$ (Stöckhert *et al.*, 1999) respectively.

between the temperatures for low-T and high-T mylonites, this does not exclude the possibility that the low-T mylonites formed after the high-T mylonite. The cooling histories deduced from the K-Ar ages, however, do not clearly show the regional temperature difference during mylonite formation due to the rapid cooling of the granitic rocks and the wide error range of the K-Ar ages.

Mylonite samples on the western side of the cataclasite zone in the northern area show younger K-Ar ages of hornblende and biotite than those in other areas. This means that the western side of the northern area has a different cooling history. This different cooling history suggests that the granitic body had intruded later than those of the other areas or that the western side in the northern area had been exhumed from deeper part. K-Ar ages of hornblende and biotite of granitic rocks do not show their intrusion age, and it is difficult to infer which possibilities are reasonable for this different cooling history. In spite of the difference of the cooling history, the deformation temperature of the high-T mylonites in the western side of the northern area is almost the same as that of the other areas (Shigematsu *et al.*, 2003). This suggests that, in either case, mylonite formation in this area occurred not at ca. 110 Ma but probably later than that in the other areas.

The fission-track ages of zircon and apatite from samples 03082105 and 03082106 are older than those of the other samples, suggesting that the areas where the low-T mylonite is developed cooled more rapidly than the other areas at a temperature lower than the closure temperature of biotite. Low-T mylonites are often accompanied by crush zones in deformation-localized zones (Tomita *et al.*, 2002). The rocks of the crush zones in deformation-localized zones have lower $\delta^{18}\text{O}$ values than the weakly deformed rocks and other typical granitic rocks in Japan, and the most probable reason for the decrease in $\delta^{18}\text{O}$ is interaction with near-surface derived water (Fujimoto *et al.*, 2002). Infiltration of near-surface derived water would accelerate cooling of granitic bodies, and this may be a possible reason why the areas where low-T mylonite is distributed cooled more rapidly than the other areas. Hornblende and K-feldspar are preserved in the low-T mylonites, suggesting that the low-

T mylonites formed under relatively dry condition (Shigematsu *et al.*, 2003) based on the P-T diagram presented by Spear *et al.* (2002). Therefore rapid cooling of the area of the low-T mylonites would occur after plastic deformation in the HFZ.

A recent geophysical study of the Nagamachi-Rifu fault, NE Japan (Nakajima *et al.*, 2001), pointed out the existence of low P -wave velocity (V_p), low S -wave velocity (V_s) and large V_p/V_s zone caused by partially molten material beneath the downward extension of the seismogenic fault. A granitic pluton cooling after intrusion would be one possible material for this partially molten zone. The results of K-Ar and fission-track dating in this study suggest that the HFZ was deformed during cooling of the granitic plutons. Iio and Kobayashi (2002) suggested the possibility that aseismic slip on the downward extensions of the seismogenic faults of intraplate earthquakes accumulate stress on the seismogenic faults and accelerate before the large intraplate earthquakes. The mylonites in the HFZ might form by plastic deformation along a downward extension of a fault during cooling of granitic plutons and would play an important role in the understanding of large intraplate earthquakes.

5. Conclusion

The K-Ar dating of the granitic rocks in the HFZ revealed that (1) the cooling history from the closure temperature of hornblende to that of biotite is almost the same along the strike of the HFZ and among granitic bodies; (2) granitic rocks in the HFZ cooled rapidly from closure temperature of hornblende to that of biotite; (3) the formation of cataclasite and mylonite occurred during the cooling of the granitic bodies; (4) the formation of mylonite and cataclasite in the HFZ started after approximately 110 Ma and terminated by 98.1 ± 2.5 Ma; (5) The cooling histories deduced from the K-Ar ages do not clearly show regional temperature differences during mylonite formation.

The fission-track ages of zircon and apatite from the granitic rocks in the HFZ imply that the areas where low-T mylonites are distributed cooled more rapidly than the other areas. This rapid cooling occurred at a temperature

lower than closure temperature of biotite after plastic deformation, and infiltration of near-surface derived water may be a possible reason why the areas where low-T mylonites are distributed cooled more rapidly than the other areas.

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